

Palynostratigraphy, palaeoclimates and palaeodepositional environments of the Miocene aged Agbada Formation in the Niger Delta, Nigeria

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ABSTRACT

A diverse assemblage of palynomorphs dominated by terrestrially derived pollen and spores is reported from three wells penetrating the Miocene Agbada Formation. The pteridophyte and bryophyte spores which form the background assemblages in the three wells are good indicators of humid tropical climates which might have prevailed in the Niger Delta during the Miocene. The abundance and variations of climate-sensitive taxa including mangrove affiliated pollen and spore types *Acrostichumsporites*, *Psilatricolporites crassus*, *Zonocostites ramonae* and *Graminidites annulatus* representing the savannah vegetation cover indicate a complex interplay between periods of wetter and drier climates.

Marine-derived dinoflagellate cysts and foraminiferal test linings are significantly present in the three wells. Taxa indicating freshwater contributions including *Botryococcus* spp., *Chomotriletes minor*, *Ovoidites parvus* and *Pediastrum* spp. are also represented numerically across the three wells.

The presence of age diagnostic palynomorphs such as *Crassoretitriletes vanraadshooveni*, *Retibrevitricholporites obodoensis*, *Tuberculodinium vancampae*, *Zonocostites ramonae* and *Tuberculodinium vancampae* recovered in the three sections studied suggest a Miocene age for the investigated Agbada Formation. The proposed age is supported by the ranges of key palynomorphs in contemporaneous basins in Africa, northern South America and other parts of the World.

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1. Introduction

The discovery of commercial quantities of hydrocarbons in Olobiri Field, Niger Delta Basin, in 1956 (Kulke, 1995) signalled the establishment of the Niger Delta as a world-class petroleum province. Since then, exploration and production activities have moved from the continent into shallow-water offshore and presently into the deep-water. This very important hydrocarbon province is considered among the world's top ten in terms of oil and gas reserves. Consequently, numerous scientific papers have been published both in international and local journals on the geology and petroleum potential of the basin. The present study looks into the application of palynology as an important tool in the determination of precise depositional settings and palaeoclimates which prevailed during deposition of some of the reservoir sequences in

the Niger Delta, specifically associated with the gross lithostratigraphic unit, the Agbada Formation (Figs. 1–3, Table 1).

Palynology finds its use in hydrocarbon exploration essentially as a stratigraphic tool in depositional settings such as continental, coastal, marginal marine and deep marine environments. When integrated with other fossil disciplines, seismic stratigraphy and mechanically acquired data including wire-line logs, the discipline can be widely applied for chronostratigraphic correlation, palaeoenvironmental studies, and evaluation of potential source, reservoir and sealing rocks (Copeland, 1993). Despite the effectiveness of palynology as an important tool in sequence stratigraphic analysis, it is still being grossly under-utilised in hydrocarbon exploration and production in low latitude marginal marine settings. This is primarily due to the limited floral turn-over of the parent vegetation types in humid, tropical low latitude vegetation belts. In cooler (high latitude) boreal settings, the additional use of marine microplankton types such as dinoflagellate cysts, psilate and ornamented acritarchs offer higher resolution

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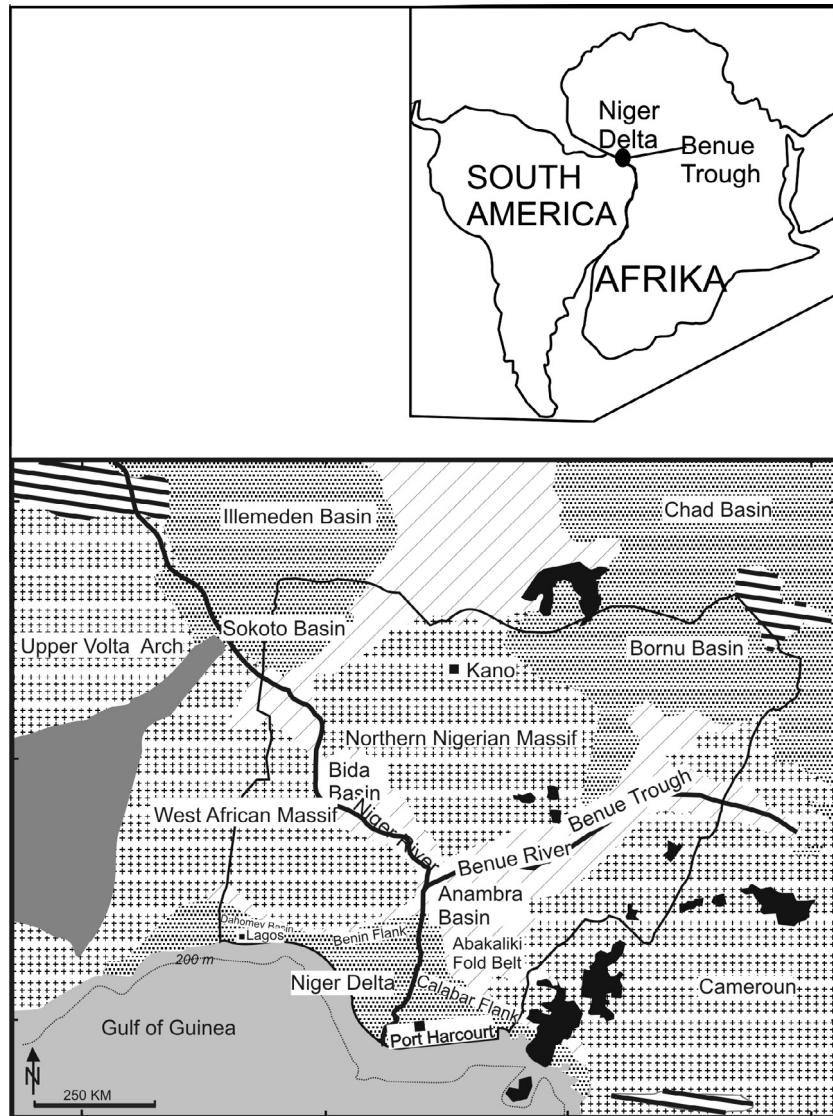


Fig. 1. Simplified geological map of Nigeria showing the location of the Niger Delta and some other Nigeria basins. Inset: Sketch map of Africa and South America (Between the Late Jurassic and Cretaceous) showing the triple junction fracture zones. Modified after the copy in Stacher (1995).

to stratigraphic subdivision being linked to more distinct oceanic climate change, namely water-mass controls and temperature.

The Cenozoic Niger Delta is located at the centre of a relict triple junction, a fracture zone which initiated the separation of continental Africa from that of South America (Fig. 1 inset) occurring between the Late Jurassic and the Cretaceous (Weber and Daukoru, 1975).

1.1. Stratigraphy of the Cenozoic Niger Delta

The lithofacies pattern of the Cenozoic Niger Delta reflects the various depositional environments peculiar to most deltas. These depositional environments, which prevailed in the sequences of the Cenozoic Niger Delta can largely be grouped into continental, transitional and marine. The Cenozoic Niger Delta covers an area of approximately 140,000 km², with cumulative sedimentary sequence of about 12,000 m (Knox and Omatsola, 1989). These sequences have been subdivided into three major sedimentary units, namely the Akata, Agbada and the Benin Formations (Table 1 and Fig. 3). The oldest of these three formations is the Eocene to Recent Akata Formation (Short and Stäuble, 1967; Reijers et al.,

1997, Table 1). The Akata Formation, characterised by continuous, uniform shale deposition was laid down in a marine environment. On top of the marine sequence is the Eocene to Recent Agbada Formation. The Agbada Formation constitutes the actual deltaic portion of the sequence. It is considered to have been accumulated in deltaic front, delta-topset and fluviodeltaic environments (Corredor et al., 2005). Capping the sequence is the mainly continental Benin Formation, deposited during the Oligocene to Recent (Reijers et al., 1997).

1.1.1. Akata Formation

The formation is characteristically composed of dark gray shale, sometimes sandy or silty of prodelta origin (Short and Stäuble, 1967; Akpoyovbike, 1978); the shales of this formation are largely undercompacted (Akpoyovbike, 1978).

1.1.2. Agbada Formation

The Agbada Formation which herein is subjected to detailed palynological investigation is composed mainly of alternating sandstone/sand bodies with mudstones. The Agbada Formation, according to Weber (1971), is considered to be composed of cyclic

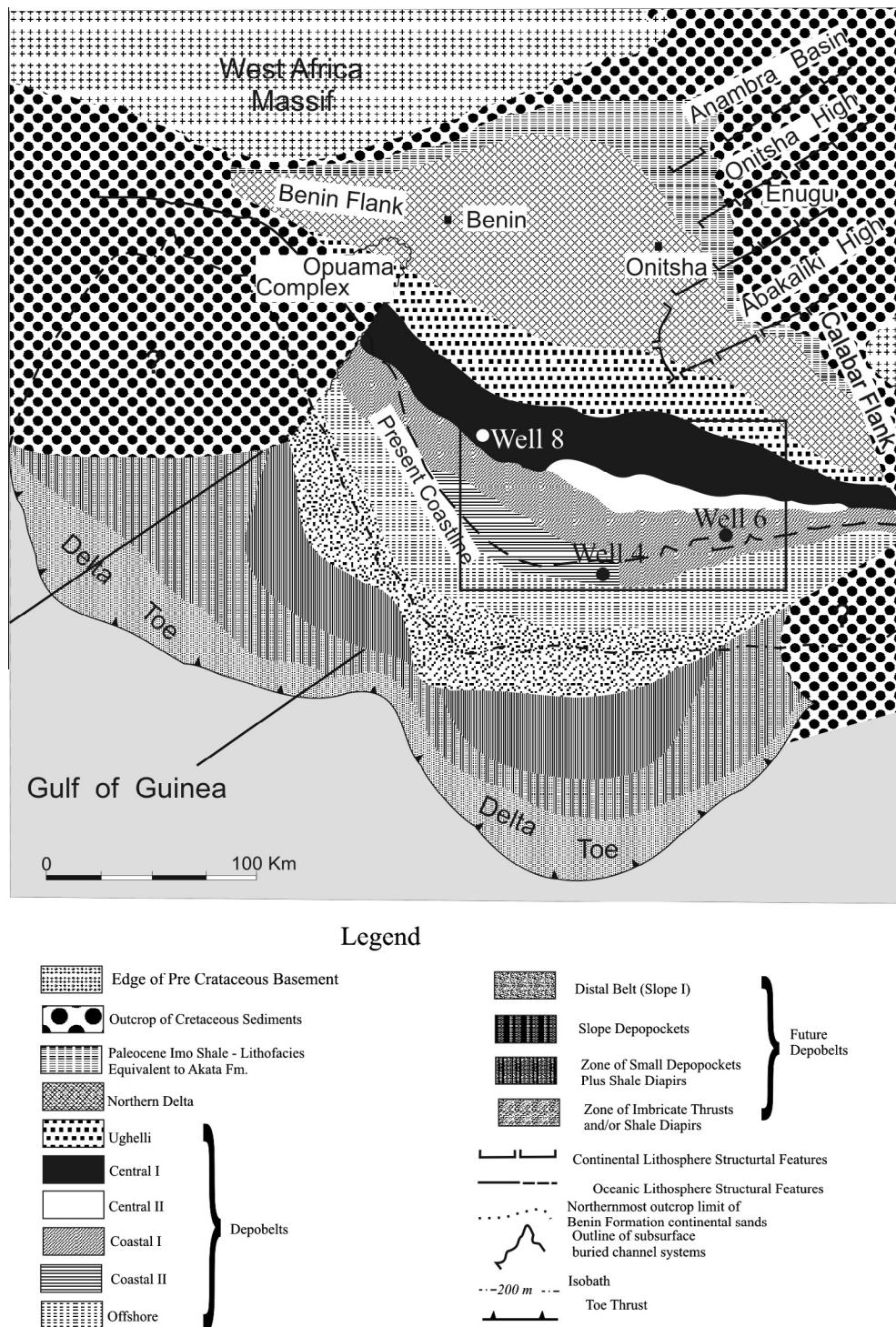


Fig. 2. Regional structural elements and depobelts of the Niger Delta. Successive phases of delta development, referred to as depobelts, may be distinguished. In each, the paralic sequence is distinct in age: for instance, the Greater Ughelli depobelト is early Miocene whereas the more seaward Central depobelト is Middle Miocene. Offshore, the delta forms two large lobes, in which depocenters are located between diapiric shale zones. Modified after Doust and Omatsola, 1990. The box indicates the location of the study area.

sequences of marine and fluvial deposits. Evidence from ditch cuttings used for the present study indicates fine to medium-grained sand/sandstones and mudstones. Lots of fairly clean coarse-grained sands have also been encountered in the samples. They are also locally calcareous, shelly, and contain pyrite. The formation as observed in the present investigation yields abundant and diverse groups of palynomorphs. Gamma Ray log data indicate high shale to sand ratio at the base of the Agbada Formation. This trend

gradually reverses itself as the middle part of the formation is approached and persists to the top. The formation grades vertically into the mainly continental deposits of the Benin Formation.

1.1.3. Benin Formation

The formation represents an almost continuous series of continental deposits (with very rare short-lived marine incursions) ranging from medium to coarse-grained sands/sandstones. Short

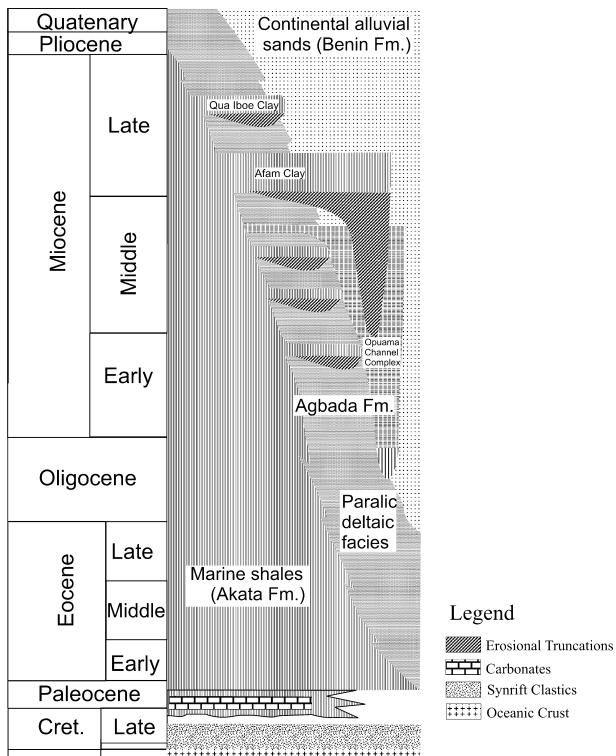


Fig. 3. Schematic representation of the diachronous nature of major lithofacies units, and the stratigraphic relationships of clay-filled channels on the delta flanks. Modified after Doust and Omatsola (1990) and Lawrence et al. (2002).

and Stäuble (1967) have described the sands/sandstones of the Benin Formation to be generally coarse-grained, very granular and pebbly but also ranging in grain-size to fine-grained. From a few samples of the Benin Formation included in the study materials used for the present investigation indicate medium to coarse-grained sands/sandstones. They are generally fairly clean but evidence of organic remnants (such as wood fragments) has been found in some of the samples.

1.1.4. Biostratigraphic outlook of the Agbada Formation

Palynological information from the Agbada Formation is often not available due to the fact that most of the palynological investigations in the Niger Delta are carried out by the oil companies and such studies are always confidential. However, biostratigraphic studies on the Agbada Formation *sensu lato* from other localities in the Niger Delta, largely based on foraminifera, have assigned an Oligocene–Pliocene age to the formation (Petters, 1979, 1983;

Ozumba, 1994; Adeniran, 1997). The same age range has also been assigned to most of the pollen recovered from the formation from other localities in the Niger Delta by Legoux (1978). This age range (Oligocene–Pliocene) points to an improvement to the earlier works of Short and Stäuble (1967) as modified after Akpoyovbiké (1978) in Table 1 herein.

1.2. Previous palynological work in the Niger Delta

A great deal of palynological investigations has been conducted in the Niger Delta. Some of the pioneering palynological investigations include those of Van Hoeken-Klinkenberg (1966) on the Maastrichtian, Paleocene and Eocene pollen and spores from Nigeria and the systematic description of new sporomorphs from the Upper Tertiary (Neogene) of Nigeria (Clarke, 1966; Clarke and Frederiksen, 1968). Shell as global pioneers in palynostratigraphy of hydrocarbon basins around the World during the 1950s and 1960s were actively collecting palynological data which was ultimately published by Germraad et al. (1968). This paper touched on the palynology of Tertiary (Cenozoic) sediments from tropical areas but focused on the Niger Delta as one of the important basins on which the publication is based. Legoux (1978), Jan du Chêne and Salami (1978) and Olofo (1992) have carried out palynological investigations including systematic description and age determination of the sedimentary sequences in the Niger Delta. The applied palynological studies in the basin include those of Morley and Richards (1993) who used Gramineae cuticle as a tool for climate change monitoring during the Late Cenozoic Niger Delta. The high resolution sequence stratigraphic work of Armentrout et al. (1999) from the Oso Field, Niger Delta, emphasised the importance of palynology in detailed sequence stratigraphic resolution. In addition, Van der Zwan and Brugman (1999), indicated the importance of palynology as one of a number of biosignals in the EA Field, Niger Delta. Ige (2009) is based on the pollen and spores record from the Niger Delta and their palaeo-vegetational implications. Oboh et al. (1992) and Oboh (1995) are other important palynological contributions in the understanding of the Niger Delta Basin.

2. Material and methods

Ditch cutting samples and log data yielded the data-base for the present investigation. The data and the location map of the selected wells were provided by the Shell Petroleum Development Company of Nigeria (SPDC).

Following standard palynological preparation method, with inputs from the procedures established at the palynological section of the Exploration Geology Department of the Technische Universität Berlin, Germany, 407 ditch cutting samples from the three

Table 1

Lithostratigraphic subdivisions of the Niger Delta, Nigeria. Modified after Akpoyovbiké (1978). The original modified after Short and Stäuble (1967).

Subsurface		Surface outcrops			
Youngest known age		Oldest known age	Youngest known age		Oldest known age
Recent	Benin Formation Afam/Qua Iboe Agbada Formation	Oligocene	Plio. Pleistocene	Benin Formation	Miocene?
Recent		Eocene	Miocene Eocene	Ogwashi-Asaba Formation Ameiki Formation	Oligocene Eocene
Recent	Akata Formation	Eocene	Late Eocene Paleocene Maestrichtian Campanian Camp./Maest. Coniacian/Santonian Turonian Albian	Imo Formation Nsuka Fm. Ajali Formation Mamu Formation Nkporo Shale Agwu Shale Eze Aku Shale Asu River Group	Paleocene Maestrichtian Maestrichtian Campanian Santonian Turonian Turonian Albian
Equivalents not known					

wells were subjected to investigation. Ten grams of each sample was washed, then treated with 10% HCl (in excess) and 35% HF for the complete removal of carbonates and the dissolution of silicates respectively. Using a 15 µm sieve, the residue was carefully sieved in an ultrasonic machine for a maximum of 5 min to improve palynomorph recovery. One to two drops of the residue are carefully dropped at the centre of each slide. A drop of warm Kaiser's Glycerol Gelatine is placed on a cover slip which is then gently lowered onto the slide containing the residue. A minimum of two slides were made from each sample.

Microscopic scanning of the palynomorphs was carried out using a Leitz Diaplan Microscope (Leitz Wetzlar; Type 307-148.001). Photographic work was done using Olympus Digital Camera DP 12. All slides, sieved and unsieved residues, are stored at the Institut für Angewandte Geowissenschaften, Technische Universität Berlin, Germany. During visual assessment of each sample, at least 200 palynomorphs (pollen, spores, dinocysts, microforaminiferal test linings, fungal haphae, freshwater/brackish marine algae and inaperturate pollen) are counted where possible. Percentage contributions of the mentioned components in each of the samples were calculated and depicted in Figs. 4–6.

3. Palynostratigraphy

Palynomorphs including pteridophyte and bryophyte spores, gymnosperm and angiosperm pollen, dinocysts, freshwater algae and fungal spores were recovered from the three wells investigated. Terrestrial plant tissues such as cuticle, xylem and woody components are also abundantly present but were not considered in the present work. Percentage quantitative composition curves, distribution patterns of the palynofloral elements encountered and range charts are shown in Figs. 4–9 and 12 respectively. Of the recovered components, terrestrially-derived forms (pollen and spores) constitute approximately 80%, whereas less than 8% are of marine origin (dinoflagellates and foraminiferal test linings) (Figs. 4–6). Freshwater algae, fungal spores and inaperturate pollen types represent the remaining 12%. The low abundance of dinoflagellates might be a consequence of a high siliciclastic discharge rate of the Proto-Niger River into the Atlantic Ocean (Gulf of Guinea) and also the tidal effects emanating from the ocean-land interface itself.

Cumulatively, 85 genera, represented by 149 palynomorph species (excluding freshwater algae, inaperturate pollen and fungi) were identified (Appendix A). Approximately 66% of the samples analysed were productive with the remaining 34% being barren.

Fern spores constitute the most abundant (ca. 42%) elements of the palynoflora recovered from the three wells (Figs. 4–6). They are present in all the samples analysed. The monolete fern spores are represented by four genera comprising the following species: *Laevigatosporites nitidus crassicooides*, *L. haardti crassicus*, *L. discordatus*, *L. javanicus*, *L. nitidus*, *Perinomonoletes* sp. 1, *P. sp.* 2, *Reticulosporites miocenicus*, *R. sp.*, *Verrucatosporites alienus*, *V. favus*, *V. balticus major*, *V. ornatus* and *V. usmensis*.

The trilete fern spores are mainly represented by: *Acrostichumsporites meghalayaensis*, *A. sp.* 1, *A. sp.* 2, *Crassoretitriletes vanraadshooveni*, *Cyathidites congoensis*, *Dictyophyllidites trilobiformis*, *Leiotriletes maxoides*, *L. microsinuosoides*, *L. spp.*, *Magnastriatites howardi*, *Todisporites flavatus*, *T. major*, *Triplanosporites microsinuosus* and *Undulatisporites structuris* among others. A comprehensive list of the species of fern spores is included in Appendix A.

Angiosperm pollen are also abundantly present in all the three studied sections. They constitute approximately 40% of the recovered palynomorphs represented herein among others by: *Prædapollis flexibilis*, *Caprifoliipites superbus*, *Crototricolpites densus*,

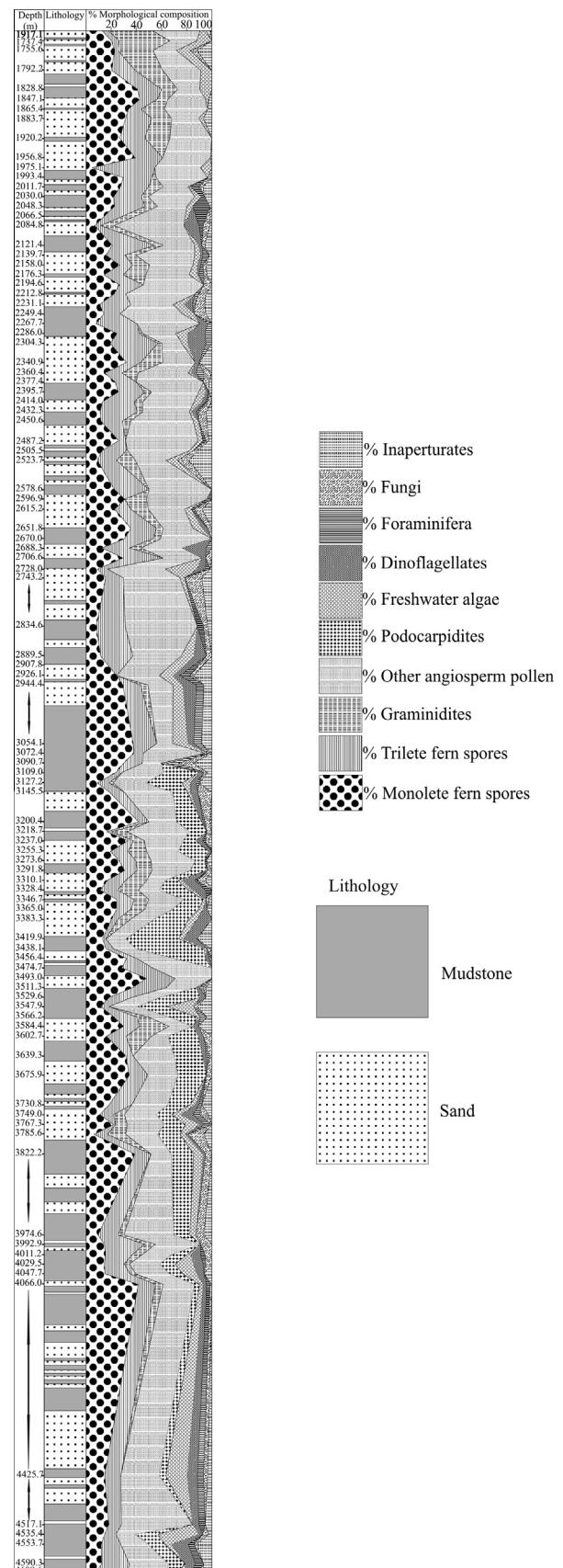


Fig. 4. Quantitative composition of the palynomorph assemblages in well 4. Note: The double headed arrows indicate large sample intervals resulting from barrenness or non-availability of samples.

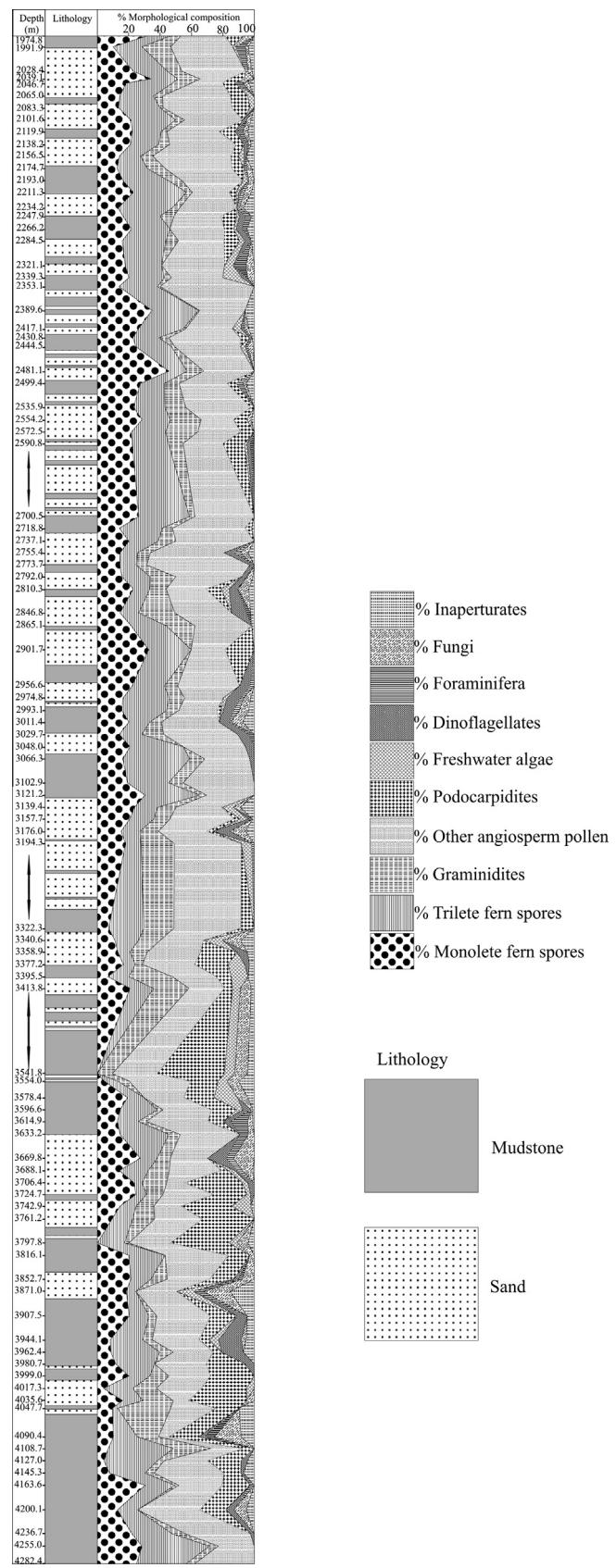
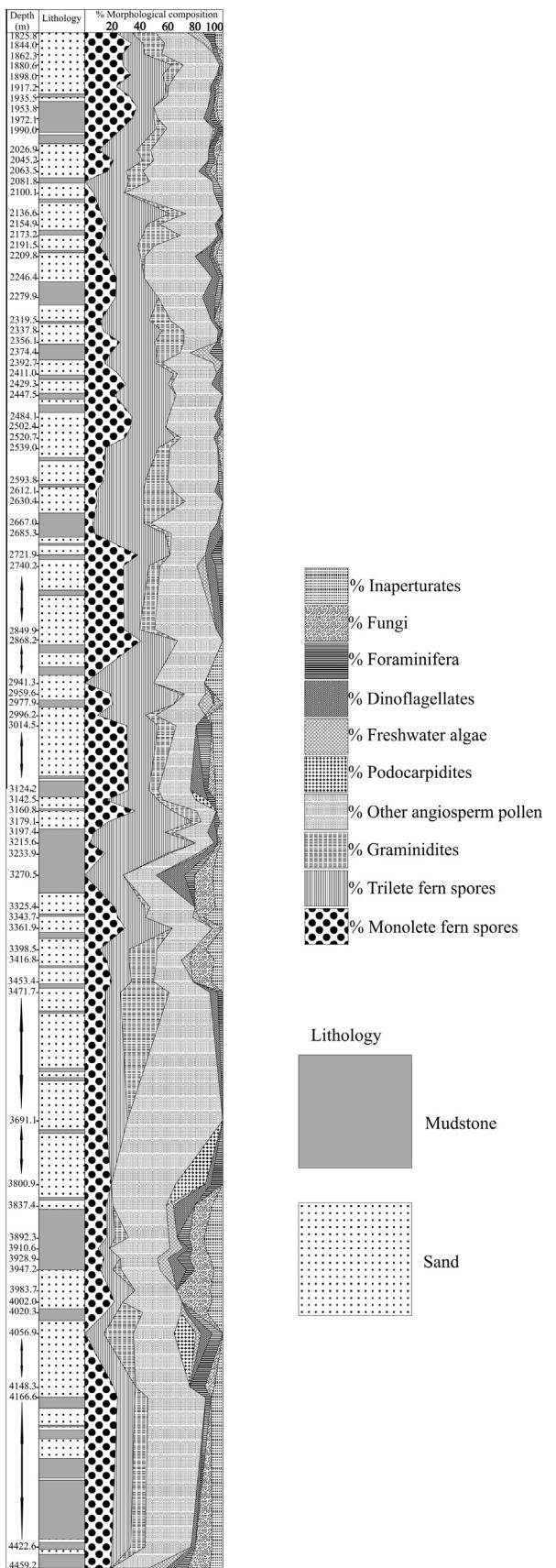


Fig. 5. Quantitative composition of the palynomorph assemblages in well 6. Note: The double headed arrows indicate large sample intervals resulting from barrenness or non-availability of samples.

Fig. 6. Quantitative composition of the palynomorph assemblages in well 8. Note: The double headed arrows indicate large sample intervals resulting from barrenness or non-availability of samples.

Ctenolophonidites costatus, *C. sp.*, *Graminidites annulatus*, *Pachydermites diederixi*, *Peregrinipollis nigericus*, *Polyadopollenites indecorus*, *Psilastephano-colporites laevigatus*, *P. cf. perforatus*, *P. sp.*, *P. spp.*, *Psilatricolporites crassus*, *Racemonocolpites hians*, *R. racematus*, *Retribrevitricolporites obodoensis*, *R. sp. cf. obodoensis*, *Retribicolporites cf. annulatus*, *R. sp. cf. guianensis*, *R. irregularis*, *Retitrescolpites cf. splendens*, *R. typicus*, *Striatopolis bellus*, *S. catatumbus*, *S. nigericus*, *S. variabilis*, *S. striatellus* and *Zonocostites cf. ramonae*. A comprehensive list of the angiosperm pollen encountered in the section is included in the Appendix A.

Another pollen group of significant abundance in the studied sections are the gymnosperms represented herein by the genera *Araucariacites* and *Podocarpidites*. The genus *Podocarpidites* is represented in large numbers in wells 4 and 8, but of limited representation in well 6 (Figs. 4–9).

Marine forms such as dinoflagellate cysts and microforaminiferal test linings are a significant component in the studied material. Dinoflagellate cysts are represented by *Lingulodinium* sp. cf. *machaerophorum* and *Lingulodinium* sp. cf. *sicula*. Other taxa with

varying degrees of representation include the genera *Adnatosphaeridium*, *Cordosphaeridium*, *Dapsilidinium*, *Homotryblium*, *Lejeuneacysta*, *Polysphaeridium*, *Spiniferites* and *Tuberculodinium*.

4. Palaeoclimate

The climate of an area is reflected by its vegetation type (Samant and Phadtare, 1997). Changes in plant communities or variations in composition/abundance of an assemblage or individual species are usually partly a direct consequence of variation in climate and/or environment. The effect of this variation on floral communities depends on whether such a climate change favours or prejudices against the plant community or individual plant in question. Qualitative and quantitative assessment of the various groups of palynomorphs undertaken here is aimed at deducing the climate phase and types of environments which prevailed during the deposition of the Agbada Formation in the Miocene.

Pteridophyte and bryophyte fern spores such as those encountered in this study including the genera *Biretisporites*,

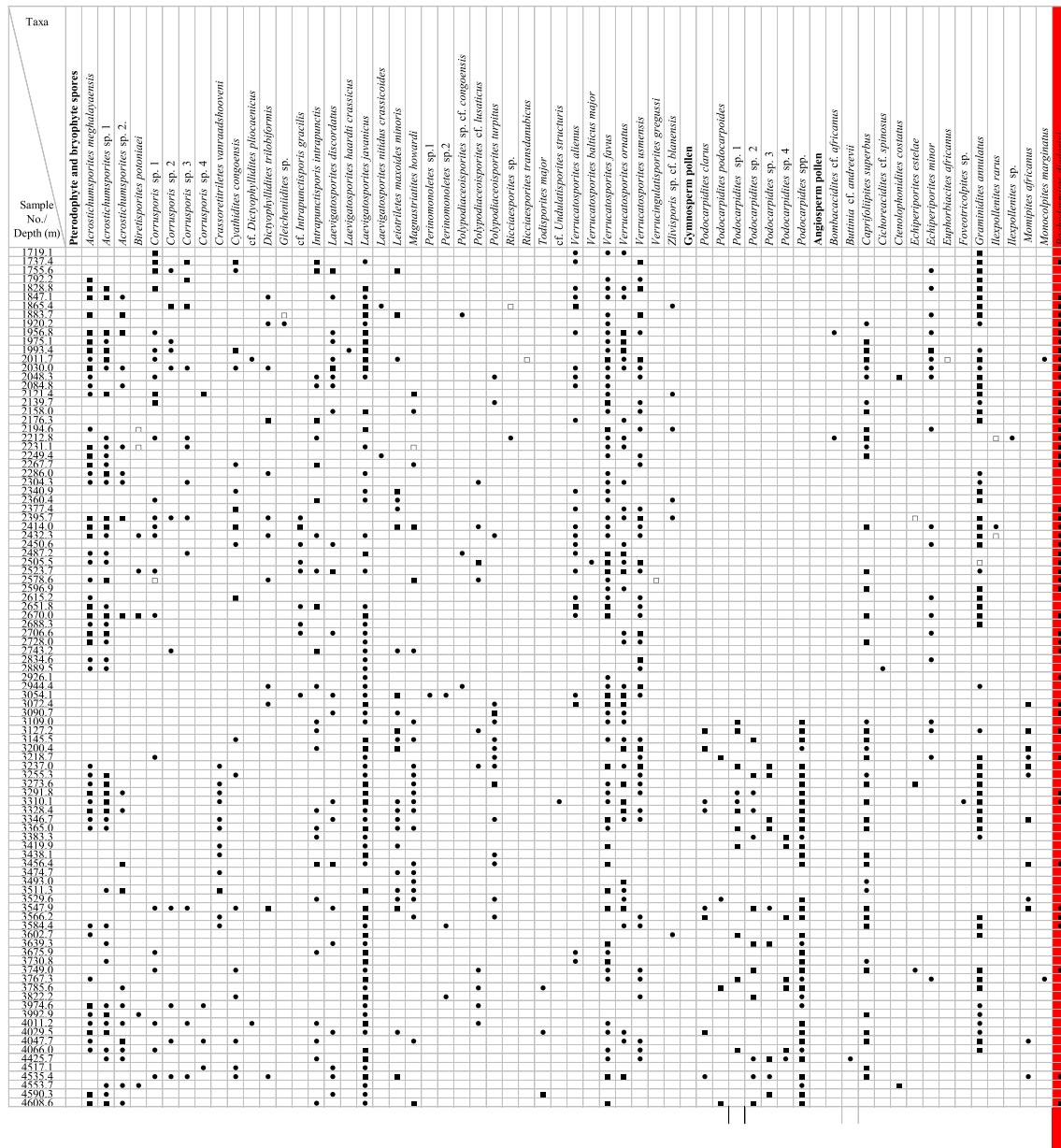


Fig. 7. Vertical distribution of palynomorphs in Well 4. Note: □ = single specimen; ● = 2–9; ■ > 10.

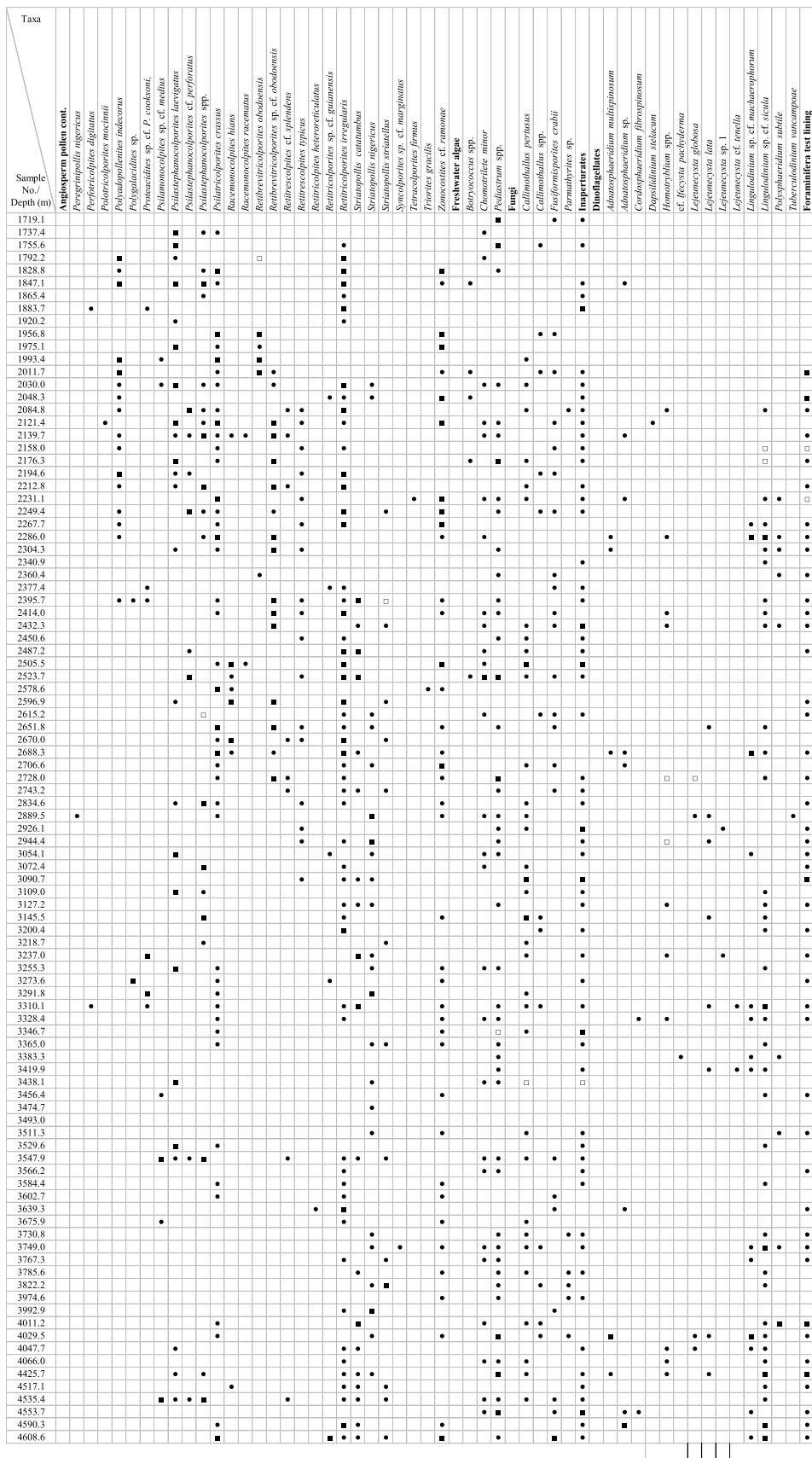


Fig. 7 (continued)

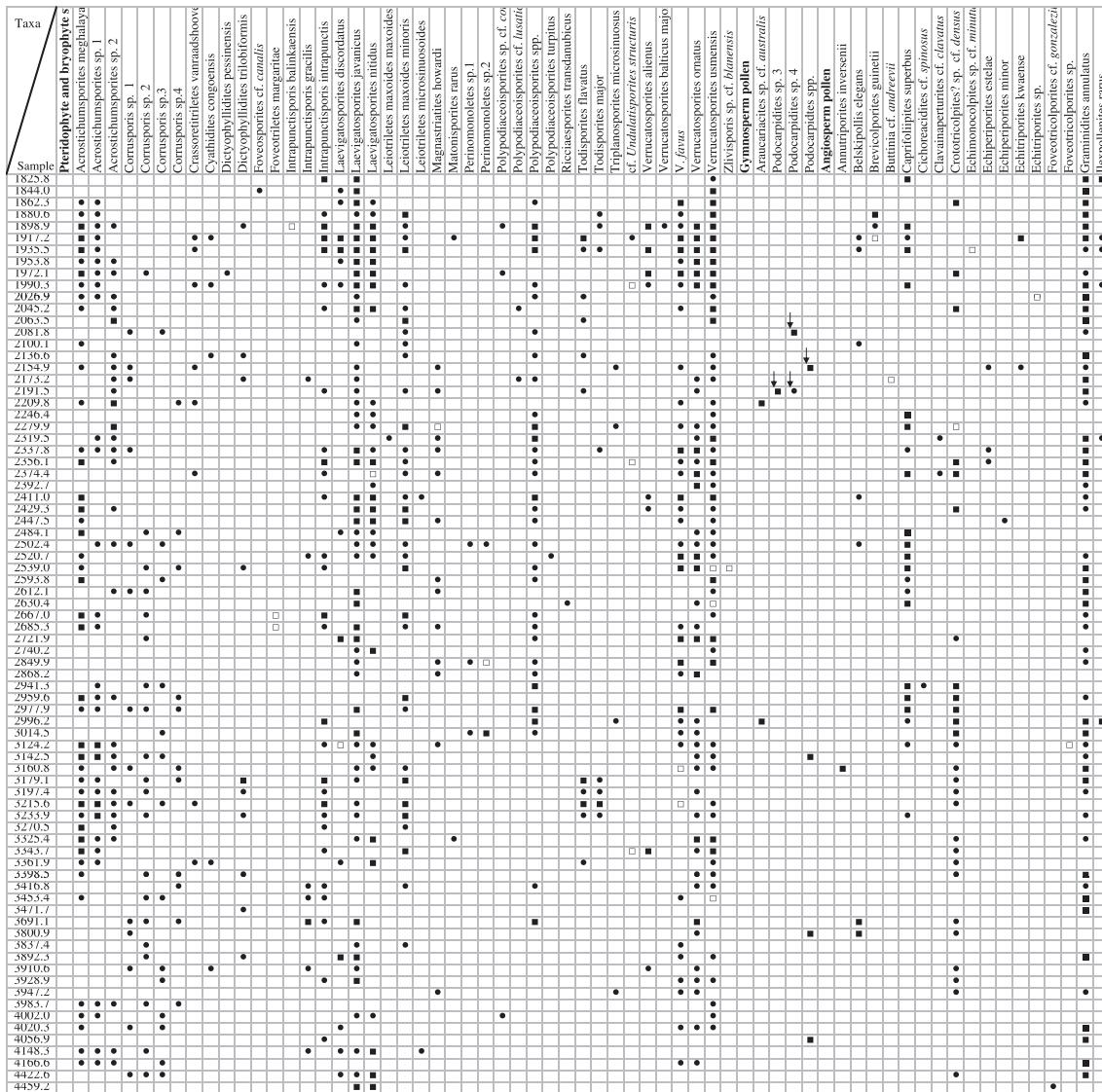


Fig. 8. Vertical distribution of palynomorphs in well 6. Note: □ = single specimen; • = 2–9; ■ = >10. Note: → = Contaminated samples resulting from sampling error.

Concavisporites, *Cyathidites*, *Dictyophylidites*, *Intrapunctisporis*, *Laevigatosporites*, *Leiotriletes*, *Magnastriatites*, *Matonispores*, *Perinomonoletes*, *Reticulosporis* and *Verrucatosporites* collectively indicate humid tropical climates (Rull, 1998, 2001). However, fluctuations are generally noticed in the palynomorphs distribution charts (Figs. 7–9) and in the quantitative composition curves (Figs. 4–6) generated from the quantitative assessment of the encountered palynomorphs. This fluctuation is a direct response of the microflora to a variety of climate, environmental and taphonomic factors witnessed within the ecosystem.

Obvious climate and environment sensitive taxa including *Acrostichumsporites meghalayaensis*, *Psilatricolporites crassus* and *Zonocostites ramonae* (mangrove taxa) and *Graminidites annulatus* (savanna to upper coastal plain grassland taxon) are abundantly recorded in all the wells investigated. These taxa have been suggested to be diagnostic of vegetation cycles attributed to climate fluctuation which could be linked to successive phases of humidity versus aridity (Armentrout et al., 1999; Poumot, 1989). The fluctuation in abundance of these taxa is directly linked to climate phase change, which in turn may be linked directly with sea-level oscillation (eustatic transgression/regression) along many other local and global factors (Poumot, 1989; Morley and Richards,

1993). Sea level changes may thus be linked with growth and loss of polar ice caps resulting in global sea level rise and fall.

As earlier indicated, mangrove taxa flourish well in prevailing warm ecosystems (average temperature greater than 22 °C) constantly under the influence of marine water inundation, a condition which prevails during sea level rise (Fig. 10) resulting in the encroachment of the coastal areas. This is also a period of increased precipitation usually associated with wetter climate conditions.

Abundant occurrence of the savanna to upper coastal plain taxon *Graminidites annulatus* is linked to a fall in sea-level (Fig. 11) associated with drier climate events. The fluctuations in abundance of these “climate” indicating taxa recorded in the present work suggest the prevalence of wetter climate conditions during relative maxima of mangrove taxa and drier climatic conditions during relative minima of mangrove taxa and vice versa for the savanna to upper coastal plain forest taxa.

A model for the flora distribution from the coastal area through the hinterland in response to climate variation resulting from sea-level oscillation is depicted in Figs. 10 and 11 (adapted from Poumot, 1989). During relative rise in sea-level (Fig. 10), there is a shift in coastline toward the continent resulting in the

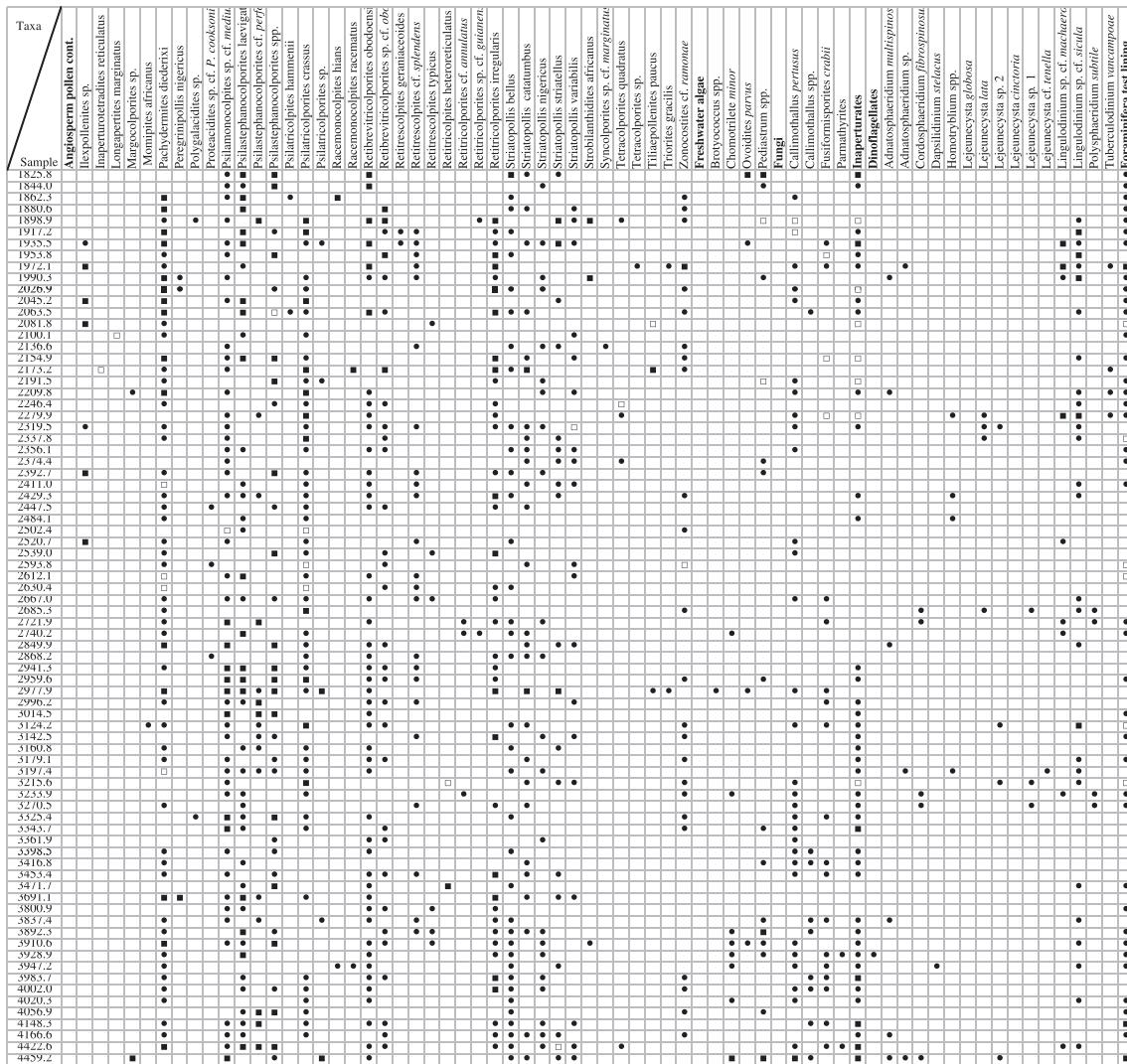


Fig. 8 (continued)

inundation of coast with marine water. During this period, the mangrove ecosystem is well established leading to the extension of mangrove vegetation or to a landward shift of the mangrove belt due to their preference for saline water. A consequence of the relative rise in sea level is also the increased precipitation leading to the prevalence of wetter climate. This is also a period where coastal vegetations such as ferns, palms and other plants typical of upper and lower coastal plains flourish. The pollen record during this period is mainly dominated by mangrove elements with some fern spores and typical marine-derived components including dinoflagellate cysts and microforaminiferal test linings. When the sea level falls (Fig. 11), the coastline is shifted in a basinward direction. The shelf area previously under marine influence is exposed and probably incised by fluvial activities. Terrestrially-derived materials are deposited in the palaeo-shelf area. This fall in sea level consequently results in a reduction in precipitation leading to the establishment of drier climate. The climate reversal ultimately results in an extension of the savanna vegetation belt (Fig. 11). The pollen record of this event is dominated by typical savanna (grass) floral elements. Note should be made in Figs. 10 and 11 of the near disappearance of savanna and mangrove ecosystem during transgressive and regressive events respectively.

4.1. Age determination

Palynological studies of the Agbada Formation elsewhere in the basin have dated the formation as Middle Miocene (Oboh, 1992, 1995). A similar study by Legoux (1978) also assigned the age range of the Agbada Formation to between Oligocene and Early Miocene.

The range chart of stratigraphically important palynomorphs recovered from the three sections is given in Fig. 12. The stratigraphically important palynomorphs selected include the following species: *Crassoretitriletes vanraadshooveni*, *Crototricolrites* sp. cf. *densus*, *Echitriporites kwaense*, *Echiperiporites minor*, *Foveotricolporites* cf. *gonzalezii*, *Retibrevitricolporites obodoensis*, *R. sp. cf. obodoensis*, *Tuberculodinium vancampoae* and *Zonocostites ramona*.

Crassoretitriletes vanraadshooveni is in general a Miocene index fossil (Germeraad et al., 1968, Nigeria, Borneo and the Caribbean; Kar and Jain, 1981, India; Lorente, 1986, Venezuela; Li and Huang, 1990, Taiwan; Oboh et al., 1992, Nigeria). The latest occurrence of the fossil according to the same authors is between Pliocene and Pleistocene. The occurrence of *Crassoretitriletes vanraadshooveni* in the studied wells is a strong indication that the deposition of the Agbada Formation occurred between the Miocene and Pleistocene.

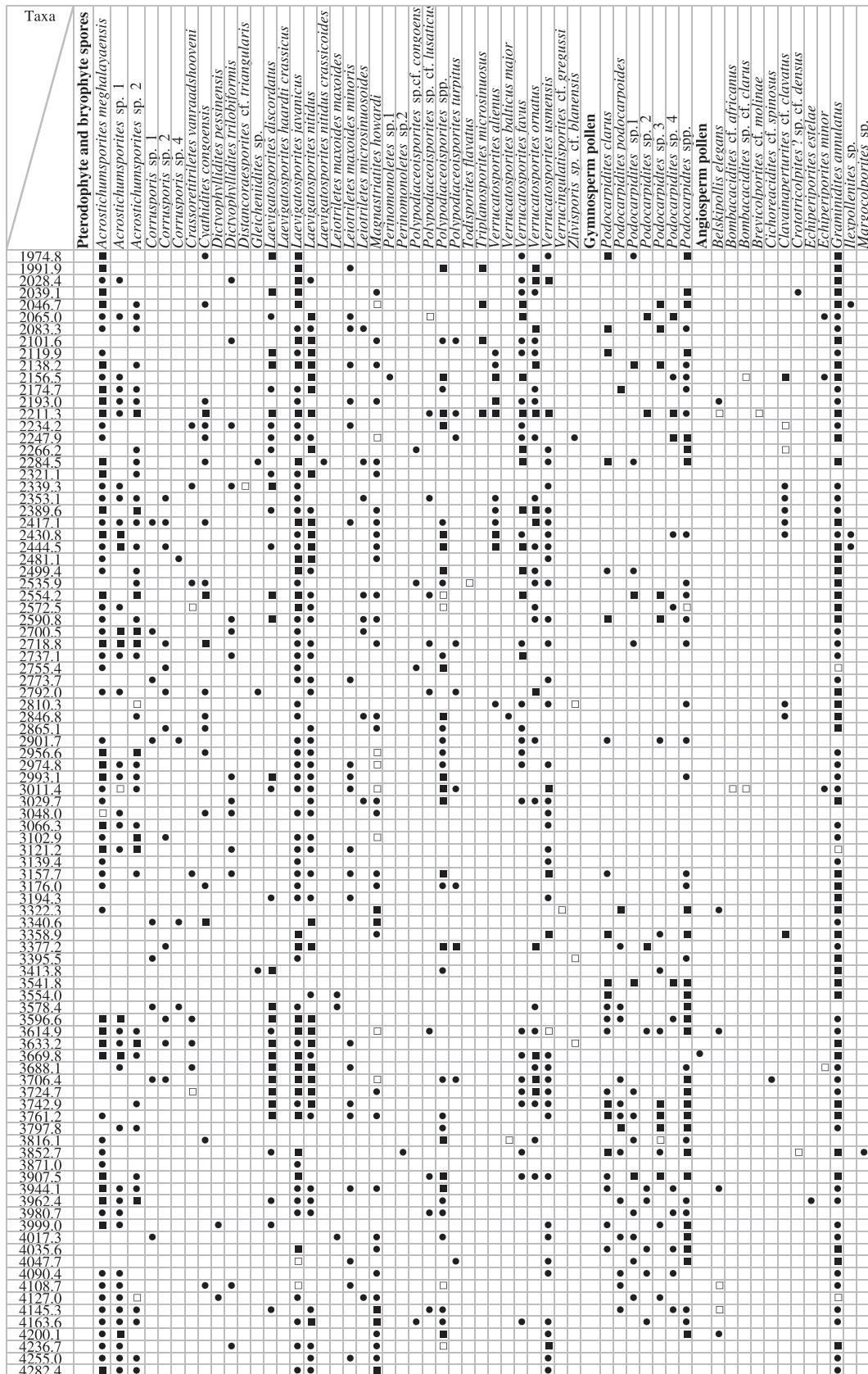


Fig. 9. Vertical distribution of palynomorphs in well 8. Note: □ = single specimen; ● = 2–9; ■ > 10.

Further, Salard-Cheboldaeff (1978, 1979) reported the ages of *Echiperiporites minor*, *Psilastephanocolporites laevigatus* and

Psilastephanocolporites perforatus to be Early Miocene. Similar age ranges have been assigned to the following: *Zonocostites ramonae*,

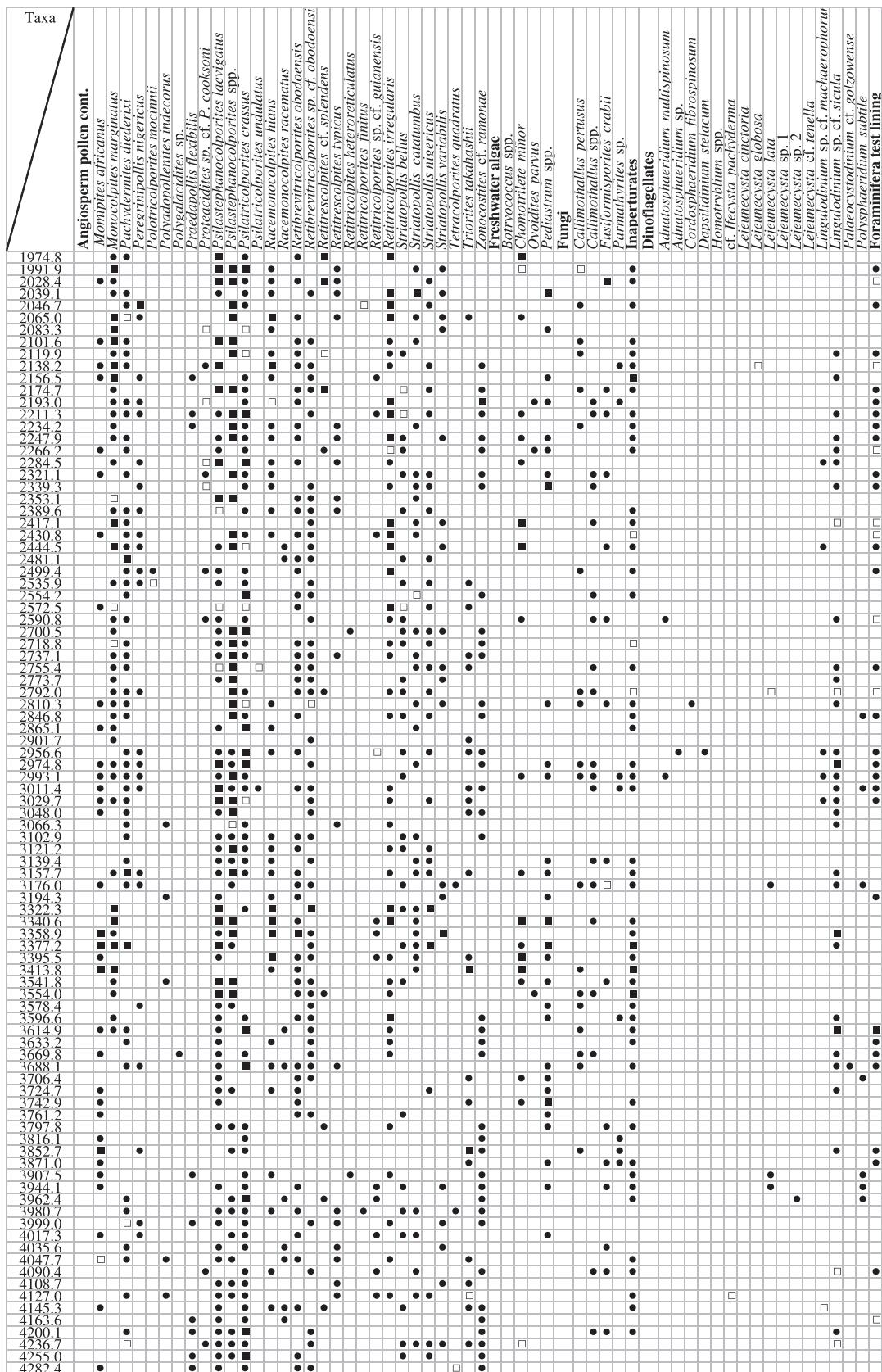


Fig. 9 (continued)

Early Miocene to Pleistocene in Nigeria (Germeraad et al., 1968); *Ricciaesporites trandanubicus*, Late Miocene (Karpasian) in

Hungary (Nagy, 1985); *Tetracolporites firmus*, *Tiliaepollenites paucus* and *Retitrescolpites geraniaceoides*, Neogene of Burundi (Sah, 1967),

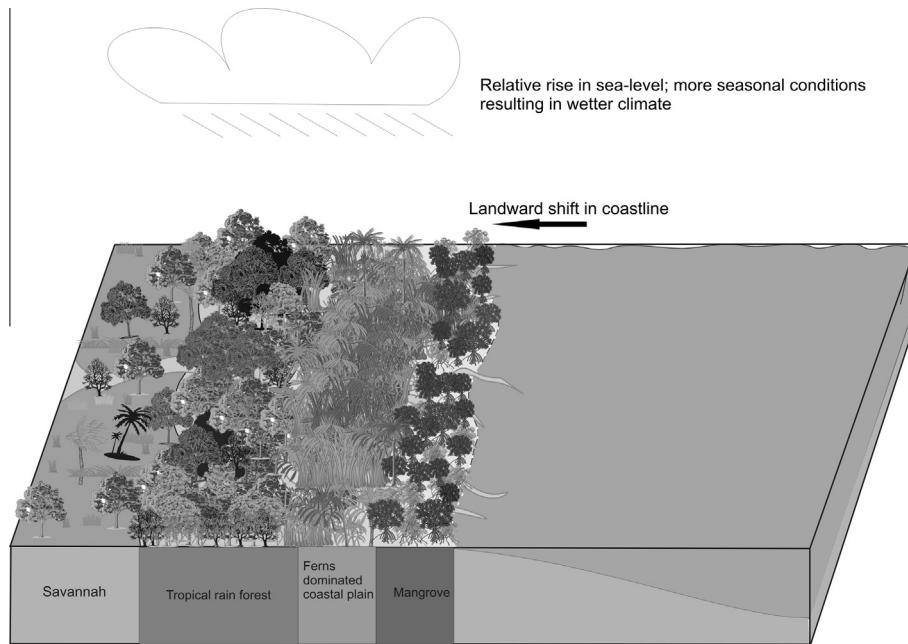


Fig. 10. Vegetation belts during relative rise in sea-level. Note: 1. The extension of the mangrove vegetation belt. 2. The decrease of the savanna belt.

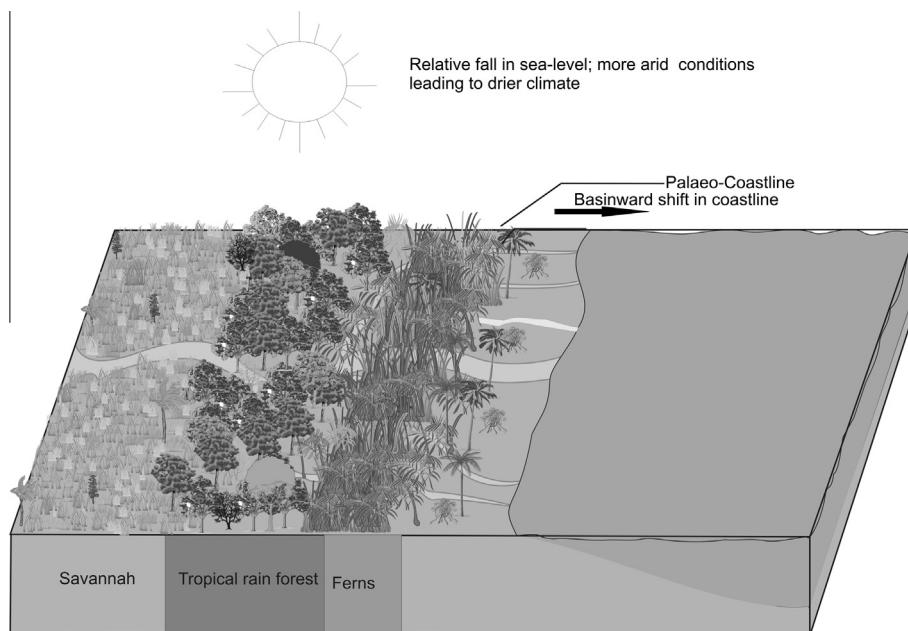


Fig. 11. Vegetation belts during the relative fall in sea-level. Note: 1. The basinward shift in coastline leading to increased fluvial activities which may include valley incision. 2. Extension of the savanna vegetation belt. 3. The near disappearance of the mangrove vegetation belt.

and *Tetracolporites quadratus*, Late Miocene to Pliocene age in Burundi and Early Miocene in Sudan (Sah, 1967 and Eisawi and Schrank, 2008 respectively). The abundant presence of the above mentioned palynomorphs at various levels in the three wells is an indication that the stratigraphic interval under investigation was deposited during the Miocene.

Palynostratigraphic evidence from marine dinocysts encountered in the studied wells largely confirms the sporomorph data cited above. Powell (1986) noted the first appearance of *Tuberculodinium vancampoae* in the early Miocene, but also mentioned conflicting earlier reports from the latest Oligocene. Originally this dinoflagellate cyst species was described from the Pleistocene

(Fensome and Williams, 2004 and sources therein). The presence of *Tuberculodinium vancampoae* in well 4 (sample 2889) and in well 6 (interval 2279.9–1972.1 m) is thus consistent with an age between Miocene and Pleistocene.

5. Discussions and conclusion

Palynological investigation of the Miocene interval of the Agbada Formation from three deep wells yielded diverse morphological groups, namely: pteridophyte and bryophyte spores, bisaccate and angiosperm pollen, dinoflagellates, foraminiferal test linings,

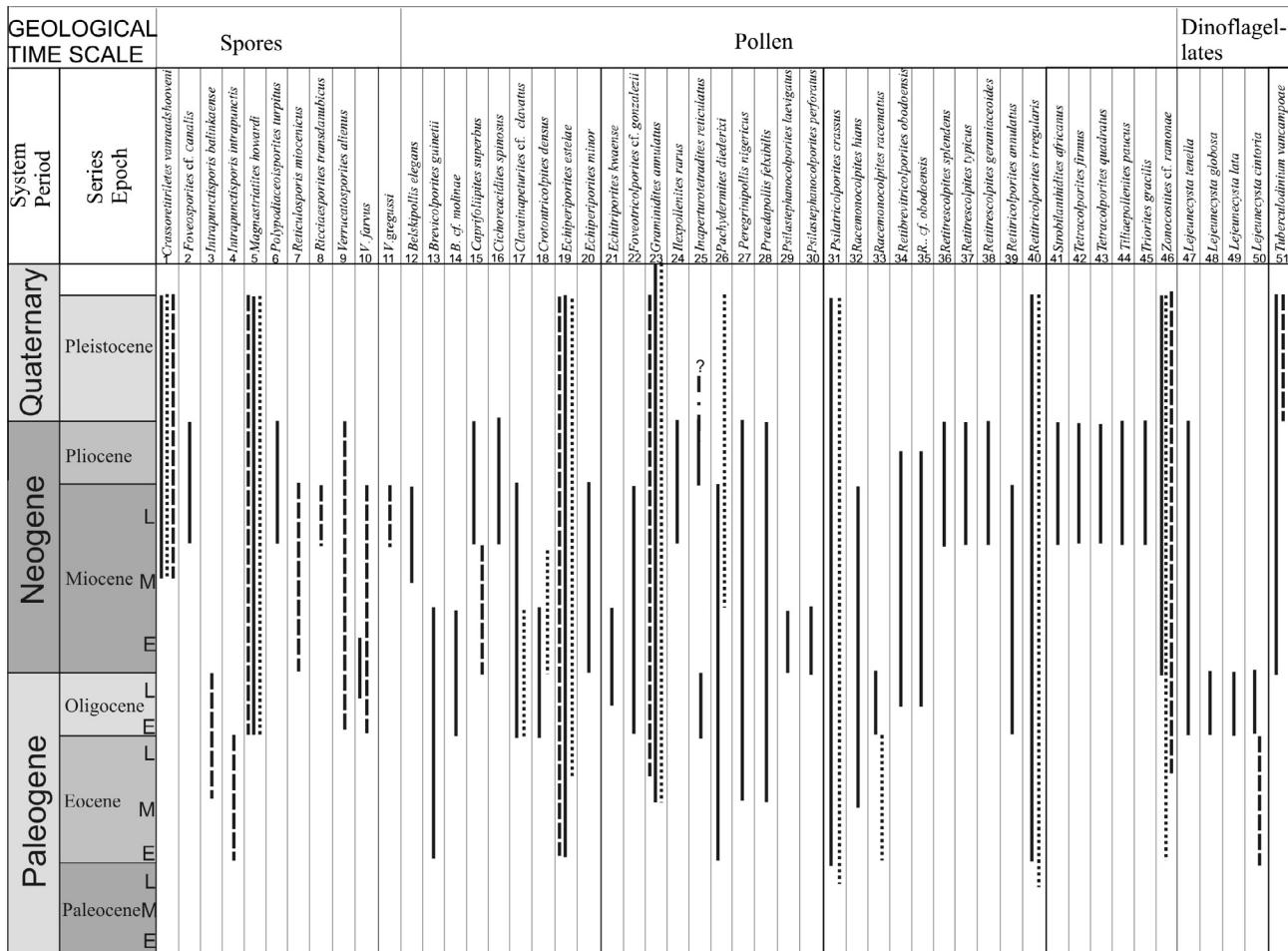


Fig. 12. Previous records of selected palynomorphs from the wells studied. Note: _____ = Ranges in Africa, = Ranges in South America, - - - - = Ranges in other parts of the world. Sources: Germeraad et al. (1968) (1, 5, 19, 23, 31, 40, 46); Lorente (1986) (1, 26, 31, 40, 46); Huang (1978) (1); Li and Huang (1990) (1); Sah (1967) (2, 6, 15, 16, 24, 36, 38, 38, 41, 42, 43, 44, 45); Eisawi and Schrank (2008) (2, 14, 17, 18, 22, 23, 26, 27, 28, 34, 36, 37, 43); Kaska (1989) (27); Clarke (1966) (27); Clarke and Frederiksen (1968) (27); Nagy (1985) (3, 8, 11); Krutzsch (1959) (4); Krutzsch (1967) (7); Takahashi and Jux (1989) (9, 10); Legoux (1978) (12, 27, 28, 32, 34, 35); Salard-Cheboldaeff et al. (1992) (12, 21, 25, 26); Salard-Cheboldaeff (1978) (13, 14, 18, 22, 25, 39); Salard-Cheboldaeff (1979) (19, 20, 21, 22, 29, 30, 31, 33); Bacchiana et al. (1982) (12, 32, 34, 35); Rao and Ramanujam (1982) (15, 18); Fasola et al. (1985) (17, 26); Obio et al. (1992) (26, 46); González-Guzmán (1967) (33); El-Beialy (1988) (47); Biffi and Grignani (1983) (47, 48, 49, 50); Bujak et al. (1980) (50); Van der Hammen and Wijnstra (1964), (17, 31, 40); Jan du Chêne and Salami (1978) (19, 40); El-Beialy (1990, 1992) (51). Species numbers 1, 5, 6, 9, 10, 16, 19, 20, 23, 26, 27, 29, 31, 32, 33, 34, 35, 40, 46, 47, 48, 49 and 50 occur in wells 4, 6 and 8; species numbers 4, 8, 13, 15, 24, 30, 45 and 51 occur in wells 4 and 6; species number 11 occurs in wells 4 and 8; species numbers 12, 17, 18, 36, 37 and 43 occur in wells 6 and 8; species number 42 occurs only in well 4; species numbers 2, 3, 7, 21, 22, 25, 38, 39, 41 and 44 occur only in well 6; species numbers 14 and 28 occur only in well 8.

freshwater algae, fungi (fungal hyphae) and inaperturate palynomorphs. Of the forms encountered, pteridophyte and bryophyte spores constitute the background assemblage in the three wells. These pteridophytes and bryophytes which include ferns are good indicators of humid tropical climates which might have prevailed in the Niger Delta since the Miocene. The abundance variations of climate-sensitive taxa including *Acrostichumsporites*, *Psilatricolporites crassus* and *Zonocostites ramonae* for the mangrove and *Graminidites annulatus* representing the savanna vegetation indicate interplay between wetter and drier climates.

Sea level oscillation is considered to play an important role in the development of a mangrove ecosystem (Vedel et al., 2006). Mangrove communities have been indicated to flourish during relative sea level rise which corresponds to a wetter climate. Abundance of the savanna taxon, *Graminidites annulatus* is suggested to be associated with a period of relative sea level fall (Poumot, 1989; Morley and Richards, 1993), indicating a possible drier climate. Rarity of cuticles and pollen of Gramineae is interpreted to be related to wetter climatic conditions (Morley, 1981; Morley and Richards, 1993). Intermediate frequencies of the two groups of climate indicators (mangrove and savanna taxa) are taken to

represent a condition between wetter and drier climates, possibly temperate conditions.

In the case of *Podocarpidites* (a gymnosperm bisaccate pollen type), its significant recovery in the sections may indicate a contribution even from distant "montane" environments, presumably from the Cameroun mountain range.

Marine-derived components represented by dinoflagellates and foraminiferal test linings are significantly present in the three wells. A relative high population of dinoflagellates in wells 6 and 4 in comparison to well 8 (Figs. 4–6, i.e. the least amount of dinoflagellates was recovered from well 8) is observed. This may suggest an increase of marine-derived components in an offshore direction as expected.

Taxa indicating freshwater contributions including *Botryococcus* spp., *Chomotriletes minor*, *Ovoidites parvus* and *Pediastrum* spp. are significantly represented across the three sections.

Well preserved palynomorphs in the studied well (Plates 1–3) suggest probably rapid burial. The absence of palynomorphs over some extensive sand intervals most especially in well 6 (Fig. 5) may suggest exposure to circulating water leading to oxidation of the palynomorphs.

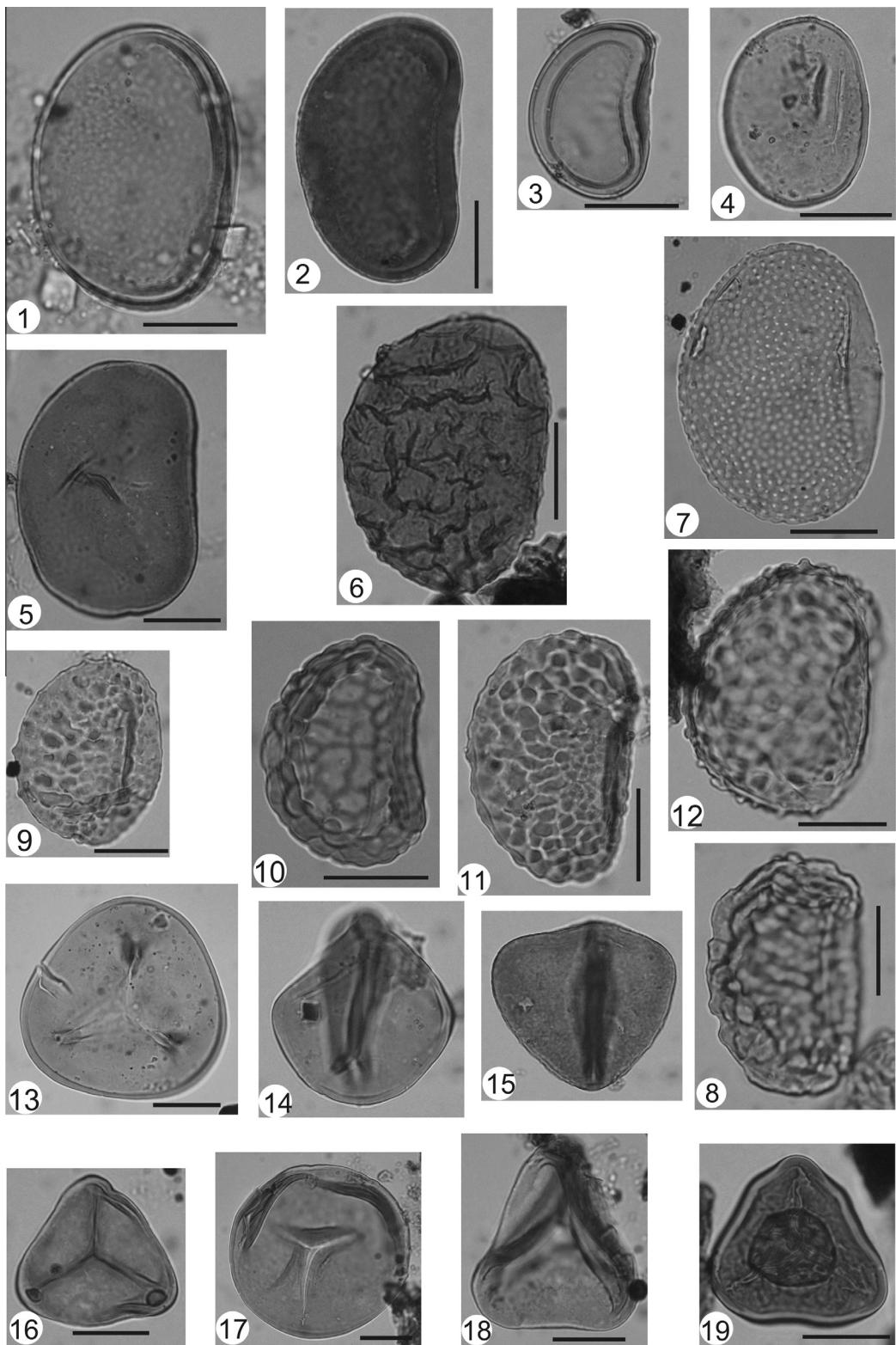


Plate 1. Fig. 1, *Laevigatosporites javanicus* TAKAHASHI, 1982. Fig. 2, *Laevigatosporites nitidus crassicooides* KRUTZSCH, 1967. Fig. 3, *Laevigatosporites haardti* (POTONIÉ & VENITZ) THOMSON & PFLUG, 1953 *crassicus* KRUTZSCH, 1967. Fig. 4, *Laevigatosporites discordatus* PFLUG, 1953. Fig. 5, *Laevigatosporites nitidus* (MAMCZAR, 1960) KRUTZSCH, 1967. Fig. 6, *Perinomonoletes* sp. 1. Fig. 7, *Reticulosporites miocenicus* (SELLING, 1944) KRUTZSCH, 1959. Fig. 8, *Verrucatosporites alienus* (POTONIÉ, 1931) THOMSON and PFLUG, 1953. Fig. 9, *Verrucatosporites favus* (POTONIÉ, 1931c) THOMSON and PFLUG, 1953 subsp. *favus*. Fig. 10, *Verrucatosporites balticus major* KRUTZSCH, 1967. Fig. 11, *Verrucatosporites ornatus* SAH, 1967. Fig. 12, *Verrucatosporites usmensis* VAN DER HAMMEN, 1956. Fig. 13, *Leiotriletes maxoides maxoides* KRUTZSCH, 1962. Fig. 14, *Leiotriletes microsinuosoides* KRUTZSCH, 1962. Fig. 15, *Triplanosporites microsinuosus* PFLANZL, 1955. Fig. 16, *Cyathidites congoensis* SAH, 1967. Fig. 17, *Todisporites major* COUPER, 1958. Fig. 18, *Dictyophyllidites trilobiformis* SAH, 1967. Fig. 19, cf. *Undulatisporites structuris* KRUTZSCH, 1962.

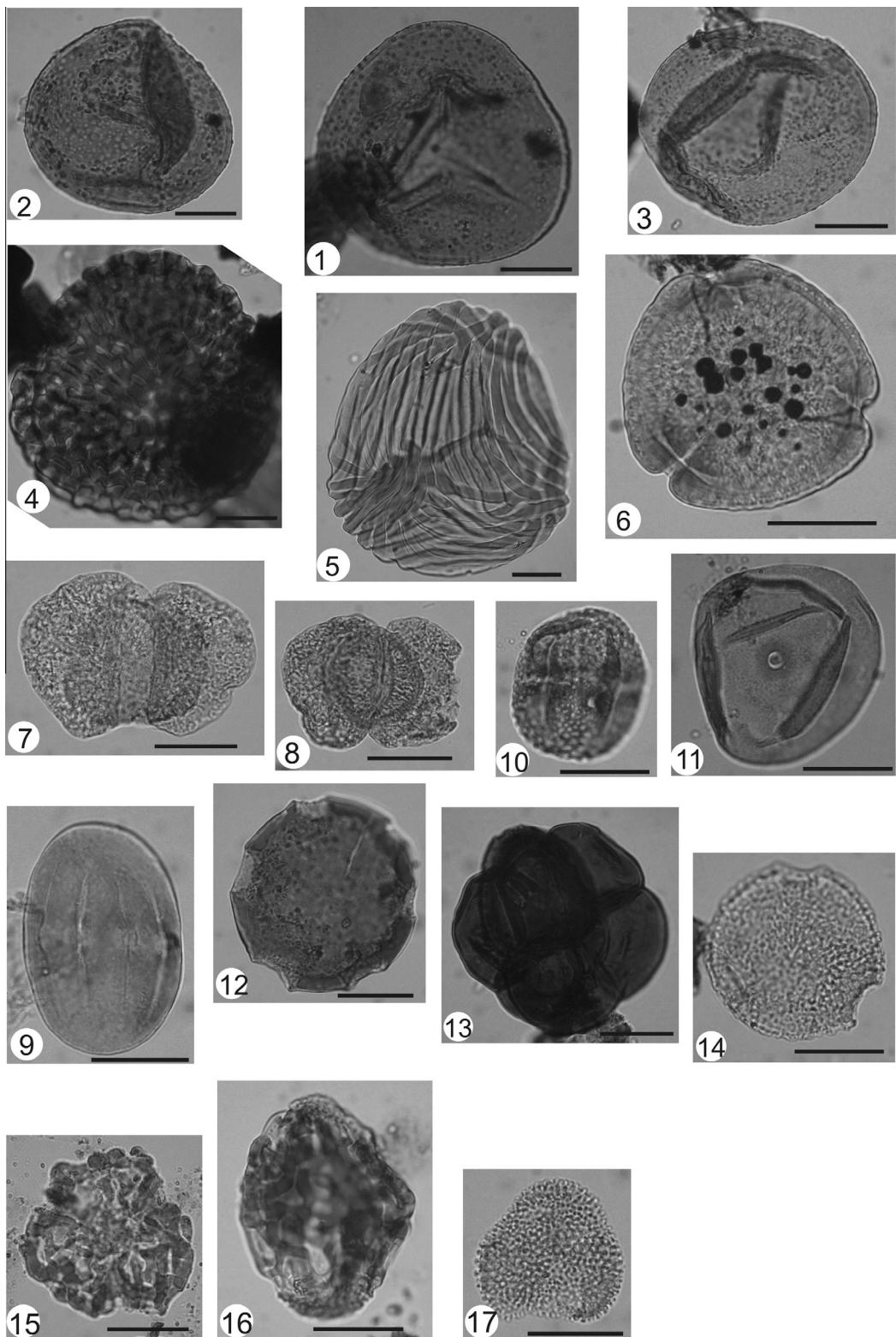


Plate 2. Fig. 1, *Acrostichumsporites meghalayaensis* KAR, 1992. Fig. 2, *Acrostichumsporites* sp. 1. Fig. 3, *Acrostichumsporites* sp. 2. Fig. 4, *Crassoretitriletes vanraadshooveni* GERMERAAD et al., 1968. Fig. 5, *Magnastriatites howardi* GERMERAAD et al. 1968. Fig. 6, *Psilatricolporites crassus* VAN DER HAMMEN and WYMSTRA, 1964. Fig. 7, *Podocarpidites clarus* SAH, 1967. Fig. 8, *Podocarpidites podocarpoides* (THIERGART, 1958) KRUTZSCH, 1971. Fig. 9, *Psilastphanocolporites laevigatus* SALARD-CHEBOLDAEFF, 1978. Fig. 10, *Psilastphanocolporites* cf. *perforatus* SALARD-CHEBOLDAEFF, 1978. Fig. 11, *Graminidites annulatus* (VAN DER HAMMEN) POTONIÉ, 1960. Fig. 12, *Pachydermites diederixi* GERMERAAD, HOPPING & MULLER, 1968. Fig. 13, *Polyadopollenites indecorus* TAKAHASHI and JUX, 1989. Fig. 14, *Caprifoliipites superbus* SAH, 1967. Figs. 15–16, *Peregrinipollis nigericus* CLARKE, 1966. Fig. 17, *Crototricolpites?* sp. cf. *densus* SALARD-CHEBOLDAEFF, 1978.

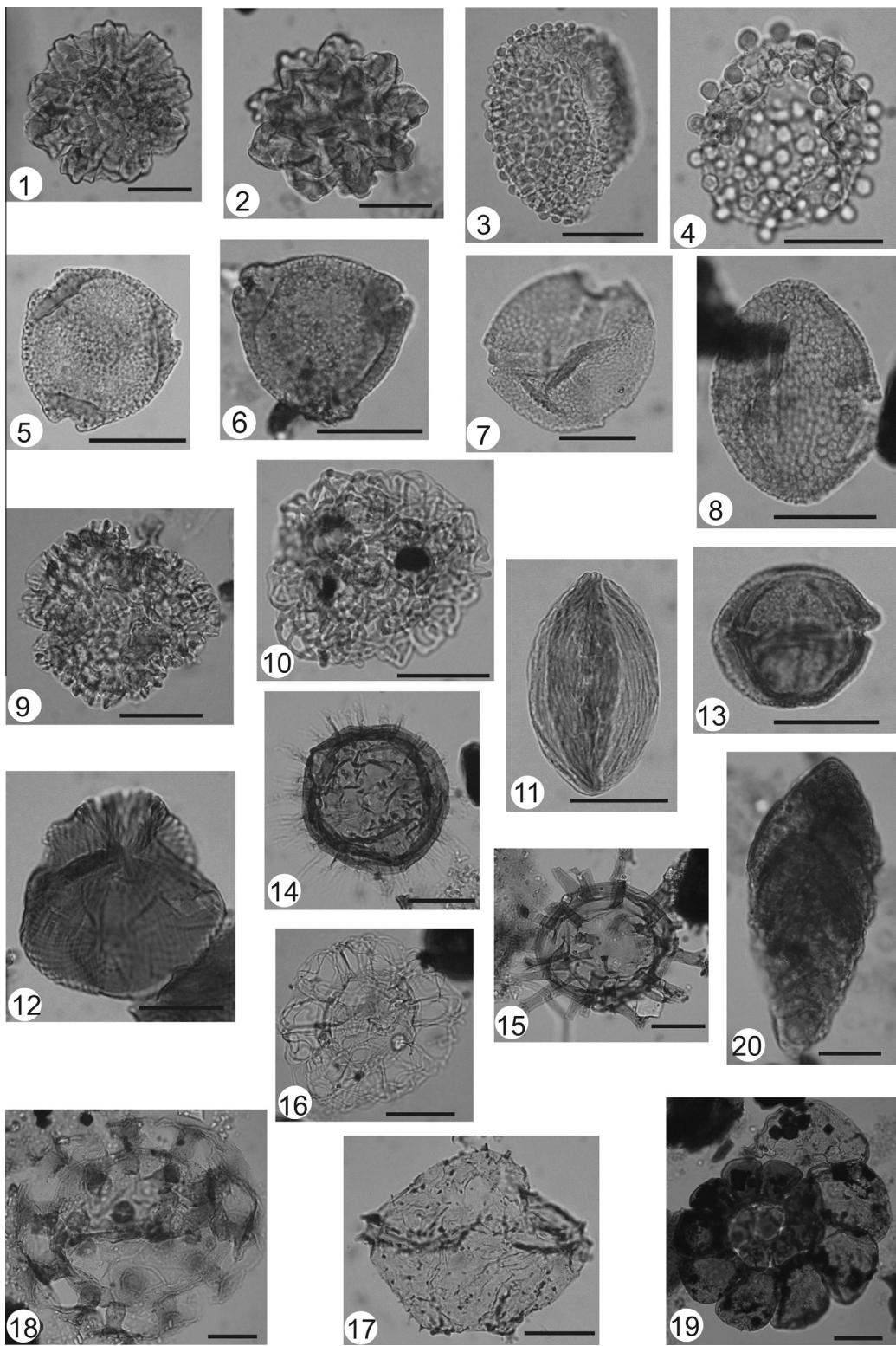


Plate 3. Fig. 1, *Ctenolophonidites costatus* VAN HOEKEN-KLIKENBERG, 1966. Fig. 2, *Ctenolophonidites* sp. Fig. 3, *Racemonocolpites hians* LEGOUX, 1978. Fig. 4, *Racemonocolpites racematus* (VAN DER HAMMEN, 1954a) GONZÁLEZ GUZMÁN, 1967. Fig. 5, *Retibrevitricolporites obodoensis* LEGOUX, 1978. Fig. 6, *Retibrevitricolporites* sp.cf. *obodoensis* LEGOUX, 1978. Fig. 7, *Retitricolporites* cf. *annulatus* SALARD-CHEBOLDAEFF, 1978. Fig. 8, *Retitricolporites* sp. cf. *guianensis* VAN DER HAMMEN and WYMSTRA, 1964. Fig. 9, *Retitrescolpites typicus* SAH, 1967. Fig. 10, *Praedapollis flexibilis* LEGOUX, 1978. Fig. 11, *Striatopollis nigericus* TAKAHASHI and JUX, 1989. Fig. 12, *Striatopollis bellus* SAH, 1967. Fig. 13, *Zonocostites* cf. *ramonae* GERMERAAD, HOPPING & MULLER, 1968. Fig. 14, *Lingulodinium* sp. cf. *machaerophorum* (Deflandre & Cookson) Wall, 1967. Fig. 15, *Dapsilidinium stelacum* Islam, 1983a. Fig. 16, *Adnatosphaeridium multispinosum* Williams & Downie, 1966c. Fig. 17, *Lejeuneacysta* cf. *tenella* (Morgenroth, 1966) Wilson and Clowes, 1980. Fig. 18, *Tuberculodinium vancampoae* (Rossignol, 1962) Wall, 1967. Figs. 19–20, Foraminifera test linings.

Appendix A. List of species

A.1. Pteridophyte and bryophyte spores

Monolete spores		Genus	<i>Distancoraesporis</i> (Krutzsch, 1963) Srivastava, 1973
Genus	<i>Laevigatosporites</i> Ibrahim, 1933. <i>Laevigatosporites discordatus</i> Pflug, 1953 <i>Laevigatosporites haardti</i> (Potonié & Venitz) Thomson & Pflug, 1953 <i>crassicus</i> Krutzsch, 1967 <i>Laevigatosporites javanicus</i> Takahashi, 1982 <i>Laevigatosporites nitidus</i> (Mamczar, 1960) Krutzsch, 1967 <i>Laevigatosporites nitidus crassicoides</i> Krutzsch, 1967	Genus	<i>Foveosporites</i> Balme, 1957 <i>Foveosporites cf. canalis</i> Balme, 1957 Genus <i>Foveotriletes</i> (Van der Hammen, 1954) ex R. Potonié, 1956 <i>Foveotriletes margaritae</i> (Van der Hammen) Germeraad, Hopping & Muller, 1968
Genus	<i>Perinomonoletes</i> Krutzsch, 1967 <i>Perinomonoletes</i> sp. 1 <i>Perinomonoletes</i> sp. 2	Genus	Genus <i>Gleicheniidites</i> Ross, 1949 <i>Gleicheniidites</i> sp. Salard-Cheboldaef, 1992
Genus	<i>Reticulosporis</i> Krutzsch 1959 <i>Reticulosporis miocenicus</i> (Selling, 1944) Krutzsch, 1959 <i>Reticulosporis</i> sp. A <i>Reticulosporis</i> sp. B	Genus	Genus <i>Intrapunctisporis</i> Krutzsch, 1959 <i>Intrapunctisporis balinkaënsis</i> Kedves, 1973 cf. <i>Intrapunctisporis gracilis</i> Krutzsch, 1967
Genus	<i>Verrucatosporites</i> Pflug & Thomson in Thomson & Pflug, 1953 <i>Verrucatosporites alienus</i> (Potonié, 1931) Thomson & Pflug, 1953 <i>Verrucatosporites balticus major</i> Krutzsch, 1967 <i>Verrucatosporites favus</i> (R. Potonié, 1931) Thomson & Pflug, 1953 subsp. <i>favus</i> <i>Verrucatosporites ornatus</i> Sah, 1967 <i>Verrucatosporites usmensis</i> Van der Hammen, 1956	Genus	<i>Intrapunctisporis intrapunctis</i> Krutzsch, 1959 Genus <i>Leiotriletes</i> (Naumova) ex Potonié & Kremp, 1954 <i>Leiotriletes maxoides</i> Kutzsch, 1962 <i>Leiotriletes maxoides minoris</i> Krutzsch, 1962 <i>Leiotriletes microsinuosoides</i> Krutzsch, 1962 L. spp.,
Trilete spores		Genus	<i>Magnastriatites</i> Germeraad, Hopping and Muller, 1968 <i>Magnastriatites howardi</i> Germeraad et al., 1968
Genus	<i>Acrostichumsporites</i> Kar, 1991 <i>Acrostichumsporites meghalayaensis</i> Kar, 1991 <i>Acrostichumsporites</i> sp.1 <i>Acrostichumsporites</i> sp.2	Genus	<i>Matonisporites</i> Couper, 1958 <i>Matonisporites rarus</i> Sah, 1967
Genus	<i>Biretisporites</i> Delcourt & Sprumont, 1955 <i>Biretisporites potoniaei</i> Delcourt & Sprumont, 1955 <i>Biretisporites</i> sp.	Genus	<i>Polypodiaceoisporites</i> Potonié, 1951 <i>Polypodiaceoisporites</i> sp. cf. <i>congoensis</i> Sah, 1967 <i>Polypodiaceoisporites</i> sp. cf. <i>lusaticus</i> Krutzsch, 1967
Genus	<i>Crassoretitriletes</i> Germeraad, Hopping and Muller, 1968 <i>Crassoretitriletes vanraadshooveni</i> Germeraad, et al., 1968 <i>Cyathidites</i> Couper, 1953	Genus	<i>Polypodiaceoisporites</i> spp. <i>Polypodiaceoisporites turpites</i> Sah, 1967
Genus	<i>Cyathidites congoensis</i> Sah, 1967 <i>Dictyophyllidites</i> Couper, 1958 <i>Dictyophyllidites pessinensis</i> (Krutzsch, 1962) Nagy, 1985 cf. <i>Dictyophyllidites plioxaenicus</i> (Thiergart, 1940) Nagy, 1985 <i>Dictyophyllidites trilobiformis</i> Sah, 1967	Genus	<i>Ricciaesporites</i> Nagy, 1968 <i>Ricciaesporites transdanubicus</i> Nagy, 1968 <i>Ricciaesporites</i> sp. <i>Todisporites</i> Couper, 1958 <i>Todisporites major</i> Couper, 1958 <i>Todisporites flavatus</i> Sah & Kar, 1969 <i>Triplanosporites</i> (Pflug) in Thomson & Pflug, 1953 <i>Triplanosporites microsinuosus</i> Pflanzl, 1955 <i>Undulatisporites</i> (Pflug) in Thomson & Pflug, 1953 cf. <i>Undulatisporites structuris</i> Krutzsch, 1962
Genus		Genus	<i>Verrucingulatisporites</i> Kedves, 1961 <i>Verrucingulatisporites</i> cf. <i>gregussi</i> Nagy, 1963
Genus		Genus	<i>Zlivisporis</i> Pacltova, 1957 <i>Zlivisporis</i> sp. cf. <i>blanensis</i> Pacltova, 1961

Alete spores		Genus	<i>Echimonocolpites</i> Van der Hammen & Garcia De Mutis, 1964
Genus	<i>Corrusporis</i> Krutzsch, 1967 <i>Corrusporis</i> sp. 1 <i>Corrusporis</i> sp. 2 <i>Corrusporis</i> sp. 3 <i>Corrusporis</i> sp. 4	Genus	<i>Echimonocolpites</i> sp. cf. <i>E. minutus</i> Salard-Cheboldaeff, 1978
		Genus	<i>Echiperiporites</i> Van der Hammen and Wymstra, 1964
			<i>Echiperiporites estelae</i> Germeraad et al., 1968
			<i>Echiperiporites minor</i> Salard-Cheboldaeff, 1975
Gymnosperm pollen		Genus	<i>Echitriporites</i> (Van der Hammen) ex. Van Hoek-Klinkenberg, 1964
Genus	<i>Araucariacites</i> (Cookson) ex Couper, 1953 <i>Araucariacites</i> sp. cf. <i>australis</i> Cookson, 1947 <i>Podocarpidites</i> (Cookson, 1947) Couper, 1953 <i>Podocarpidites clarus</i> Sah, 1967 <i>Podocarpidites podocarpoides</i> (Thiergart, 1958) Krutzsch, 1971 <i>Podocarpidites</i> sp. 1 <i>Podocarpidites</i> sp. 2 <i>Podocarpidites</i> sp. 3 <i>Podocarpidites</i> sp. 4 <i>Podocarpidites</i> spp	Genus	<i>Echitriporites kwaense</i> Salard-Cheboldaeff, 1974
		Genus	<i>Echitriporites</i> sp.
			<i>Euphorbiacites</i> (Zaklinskaja, 1965) ex. Sung & Lee, 1976
			<i>Euphorbiacites africanus</i> Takahashi and Jux, 1989
		Genus	<i>Foveotricolpites</i> Van der Hammen et Garcia, 1966
			<i>Foveotricolpites</i> sp.
		Genus	<i>Foveotricolporites</i> Pierce, 1961
			<i>Foveotricolporites</i> cf. <i>gonzalezi</i> Salard-Cheboldaeff, 1978
			<i>Foveotricolporites</i> sp.
		Genus	<i>Graminidites</i> (Van der Hammen) Potonié, 1960
			<i>Graminidites annulatus</i> (Van der Hammen) Potonié, 1960
Angiosperm pollen		Genus	<i>Ilexpollenites</i> Thiergart 1937
Genus	<i>Annutrisporites</i> González Guzmán, 1967 <i>Anntriporites iversenii</i> González Guzmán, 1967 Genus <i>Belskipollis</i> Legoux, 1978 <i>Belskipollis elegans</i> Legoux, 1978 Genus <i>Bombacacidites</i> Couper, 1960 <i>Bombacacidites</i> cf. <i>africanus</i> Takahashi and Jux, 1989 <i>Bombacacidites</i> sp. cf. <i>B. clarus</i> Sah, 1967 <i>Brevicolporites</i> Anderson, 1960 <i>Brevicolporites guinetii</i> Salard-Cheboldaeff, 1978 <i>Brevicolporites</i> cf. <i>molinae</i> (Schuler & Doubinger) Salard-Cheboldaeff, 1978 <i>Buttinia</i> Boltenhagen, 1967 <i>Buttinia</i> cf. <i>andreevii</i> Boltenhagen, 1967 <i>Caprifoliipites</i> Wodehouse, 1933 <i>Caprifoliipites superbus</i> Sah, 1967 <i>Cichoreacidites</i> Sah, 1967 <i>Cichoreacidites</i> cf. <i>spinosis</i> Sah, 1967 <i>Clavainaperturites</i> Van der Hammen and Wymstra, 1964 <i>Clavainaperturites</i> cf. <i>clavatus</i> Van der Hammen and Wymstra, 1964 <i>Crototricolpites</i> (Van Hoeken-Klinkenberg, 1964) Van Hoeken-Klinkenberg, 1966 <i>Crototricolpites?</i> sp. cf. <i>densus</i> Salard-Cheboldaeff, 1978 <i>Ctenolophonidites</i> Van Hoeken-Klinkenberg, 1966 <i>Ctenolophonidites costatus</i> Van Hoeken-Klinkenberg, 1966 <i>Ctenolophonidites</i> sp.	Genus	<i>Ilexpollenites</i> <i>rarus</i> Sah, 1967
			<i>Ilexpollenites</i> sp.
			<i>Inaperturotetradiites</i> Van Hoeken-Klinkenberg, 1964
			<i>Inaperturotetradiites reticulatus</i> Salard-Cheboldaeff, 1978
			<i>Longapertites</i> Van Hoeken-Klinkenberg, 1964
			<i>Longapertites marginatus</i> Van Hoeken-Klinkenberg, 1964
			<i>Margocolporites</i> Ramnujam, 1966
			<i>Margocolporites</i> sp. Jaramillo et al. 2007
			<i>Momipites</i> Wodehouse, 1932
			<i>Momipites africanus</i> Van Hoeken-Klinkenberg, 1966
			<i>Monocolpites</i> Van der Hammen, 1954
			<i>Monocolpites marginatus</i> Van der Hammen, 1954
			<i>Pachydermites</i> Germeraad, Hopping & Muller, 1968
			<i>Pachydermites diederixi</i> Germeraad, Hopping & Muller, 1968
			<i>Peregrinipollis</i> Clarke, 1966
			<i>Peregrinipollis nigericus</i> Clarke, 1966
			<i>Perfotricolpites</i> González Guzmán, 1967
			<i>Perfotricolpites digitatus</i> González Guzmán 1967
			<i>Polotricolporites</i> González Guzmán, 1967
			<i>Polotricolporites mocinnii</i> González Guzmán, 1967
			<i>Polyadopollenites</i> Pflug & Thomson, 1953
			<i>Polyadopollenites indecorus</i> Takahashi

(continued on next page)

Genus	and Jux, 1989 <i>Polygalacidites</i> Sah & Dutta, 1966 <i>Polygalacidites</i> sp. Takahashi and Jux, 1989	1989 <i>Striatopollis striatellus</i> (Takahashi, 1961) Takahashi in Takahashi and Kim, 1979 <i>Striatopollis variabilis</i> Takahashi and Jux, 1989
Genus	<i>Praedapollis</i> Boltenhagen and Salard, 1973	Genus <i>Strobilanthidites</i> Sah, 1967
Genus	<i>Praedapollis flexibilis</i> Legoux, 1978 <i>Proteacidites</i> (Cookson ex Couper, 1953) Martin and Harris, 1974 <i>Proteacidites</i> sp. cf. <i>P. cooksoni</i> Salard-Cheboldaeff, 1978	Genus <i>Strobilanthidites africanus</i> Sah, 1967 <i>Syncolporites</i> Van der Hammen, 1954 <i>Syncolporites</i> sp. cf. <i>marginatus</i> Van Hoeken-Klinkenberg, 1964
Genus	<i>Psilamonocolpites</i> Puri, 1963 <i>Psilamonocolpites</i> sp. cf. <i>medius</i> (Van der Hammen) Van der Hammen & García De Mutis, 1966	Genus <i>Tetracolporites</i> Couper, 1953 <i>Tetracolporites firmus</i> Sah, 1967 <i>Tetracolporites quadratus</i> Sah, 1967 <i>Tetracolporites</i> sp. Eisawi and Schrank, 2008
Genus	<i>Psilastephanocolporites</i> Leidelmeyer, 1966 <i>Psilastephanocolporites laevigatus</i> Salard-Cheboldaeff, 1978 <i>Psilastephanocolporites</i> sp. cf. <i>perforatus</i> Salard-Cheboldaeff, 1978 <i>Psilastephanocolporites</i> spp.	Genus <i>Tiliaepollenites</i> (Potoniè) Potoniè & Venitz, 1934 <i>Tiliaepollenites paucus</i> Sah, 1967
Genus	<i>Psilatricolpites</i> Van der Hammen and Wymstra, 1964 <i>Psilatricolpites hammenii</i> Boltenhagen, 1976	Genus <i>Triorites</i> (Ross) ex. Couper, 1953 <i>Triorites takahashii</i> Schrank in Bankole et al., 2007
Genus	<i>Psilatricolporites</i> (Van der Hammen) ex Pierce, 1961 <i>Psilatricolporites crassus</i> Van der Hammen and Wymstra, 1964 <i>Psilatricolporites</i> sp.	Genus <i>Triorites gracilis</i> Sah, 1967 <i>Zonocostites</i> Germeraad, Hopping and Muller, 1968
Genus	<i>Racemonocolpites</i> Gonzalez Guzman, 1967 <i>Racemonocolpites hians</i> Legoux, 1978 <i>Racemonocolpites racematus</i> (Van der Hammen, 1954a) González Guzmán, 1967 Genus <i>Retibrevitricolporites</i> , Legoux, 1978 <i>Retibrevitricolporites obodoensis</i> Legoux, 1978 <i>Retibrevitricolporites</i> sp. cf. <i>obodoensis</i> Legoux, 1978	Genus <i>Zonocostites cf. ramonae</i> Germeraad, Hopping and Muller, 1968
Genus	<i>Retitricolpites</i> Van der Hammen and Wymstra, 1964 <i>Retitricolpites heteroreticulatus</i> Boltenhagen, 1976	
Genus	<i>Retitrescolpites</i> Sah, 1967 <i>Retitrescolpites</i> cf. <i>geraniaceoides</i> Sah, 1967 <i>Retitrescolpites</i> cf. <i>splendens</i> Sah, 1967 <i>Retitrescolpites</i> typicus Sah, 1967	
Genus	<i>Retitricolporites</i> (Van der Hammen, 1956) ex. Van der Hammen and Wymstra, 1964 <i>Retitricolporites</i> cf. <i>annulatus</i> Salard-Cheboldaeff, 1978 <i>Retitricolporites</i> sp. cf. <i>gianensis</i> Van der Hammen and Wymstra, 1964 <i>Retitricolporites irregularis</i> Van der Hammen and Wymstra, 1964	
Genus	<i>Striatopollis</i> Krutzsch, 1959 <i>Striatopollis bellus</i> Sah, 1967 <i>Striatopollis catatumbus</i> (González Guzmán, 1967) Takahashi and Jux, 1989 <i>Striatopollis nigericus</i> Takahashi and Jux,	

A.2. Dinoflagellates

Gonyaulacoids	<i>Adnatosphaeridium</i> Williams & Downie, 1966c
Genus	<i>Adnatosphaeridium multispinosum</i> Williams & Downie, 1966c
Genus	<i>Adnatosphaeridium</i> sp.
Genus	<i>Cordosphaeridium</i> Eisenack, 1963
Genus	<i>Cordosphaeridium fibrospinosum</i> Davey & Williams, 1966
Genus	<i>Dapsilidinium</i> Bujak, Downie, Eaton & Williams, 1980c
Genus	<i>Dapsilidinium stelacum</i> Islam, 1983a
Genus	<i>Homotryblium</i> Davey and Williams, 1966b
Genus	<i>Homotryblium</i> spp.
Genus	<i>Ifecysta</i> Jan Du Chêne & Adediran, 1985
Genus	cf. <i>Ifecysta pachyderma</i> Jan Du Chêne & Adediran, 1985
Genus	<i>Lingulodinium</i> Wall, 1967
Genus	<i>Lingulodinium</i> sp. cf. <i>machaerophorum</i> (Deflandre & Cookson) Wall, 1967
Genus	<i>Lingulodinium</i> sp. cf. <i>sicula</i> (Drugg, 1970) Wall & Dale in Wall et al., 1973
Genus	<i>Polysphaeridium</i> Davey & Williams, 1966
Genus	<i>Polysphaeridium</i> <i>subtile</i> Davey & Williams, 1966 emend. Bujak et al., 1980
Genus	<i>Spiniferites</i>
Genus	<i>Spiniferites</i> spp.
Genus	<i>Tuberculodinium</i> Wall, 1967
	<i>Tuberculodinium vancampoae</i> (Rossignol, 1962) Wall, 1967

Peridinioids		Doust, H., Omatsola, E., 1990. Niger Delta. In: Edwards, J.D., Santogrossi, P.A. (Eds.), Divergent/pассив Margin Basins, AAPG Memoir 48: Tulsa. American Association of Petroleum Geologists, pp. 201–238.
Genus	<i>Lejeuneucysta</i> Artzner & Dörhöfer, 1978	Eisawi, A., Schrank, E., 2008. Upper Cretaceous to Neogene palynology of the Melut Basin, southeast Sudan. Palynology 32, 101–129.
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	<i>Lejeuneucysta lata</i> Biffi and Grignani, 1983	El-Beialy, S.Y., 1992. Miocene and Pliocene dinoflagellate cysts and other palynomorphs from the Damanhur South-1 well, western Nile Delta, Egypt. N. Jb. Geol. Paläontol. Mh. 1992 (10), 577–594.
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	<i>Lejeuneucysta</i> sp. 2	Fensome, R.A., Williams, G.L., 2004. The Lenten and Williams index of fossil dinoflagellates, 2004 ed. AASP Contributions Series 42, American Association of Stratigraphic Palynologists Foundation, 909pp.
	<i>Lejeuneucysta cf. tenella</i> (Morgenroth, 1966)	Germraad, J.H., Hopping, C.A., Muller, J., 1968. Palynology of Tertiary sediments from tropical areas. Rev. Palaeobot. Palynol. 6, 189–348.
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Genus	<i>Palaeocystodinium</i> Alberti, 1961	Huang, T.-C., 1978. Miocene palynomorphs of Taiwan (III) spores. Taiwania 23, 7–55.
	<i>Palaeocystodinium cf. golzowense</i> Alberti, 1961	Ige, O.E., 2009. A Late Tertiary pollen record from Niger Delta, Nigeria. Int. J. Bot. 1–13, ISSN 1811-9700.
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Genus	<i>Botryococcus</i>	Kar, R.K., Jain, K.P., 1981. Palynology of Neogene sediments around Quilon and Varkala-Kerala coast, South India – 2. Spores and pollen grains. Palaeobotanist 27, 113–131.
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